

3. SILICEOUS MICROFOSSILS IN UPPER NEogene SEDIMENTS FROM FIVE GUYOTS IN THE WESTERN NORTH PACIFIC, WITH SPECIAL EMPHASIS ON DIATOMS¹

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ABSTRACT

Upper Neogene sediments from six holes (Holes 871A, 872A, 872C, 873B, 878A, and 880A) drilled on five guyots during Leg 144 along a south-north transect were quantitatively analyzed for the abundance of siliceous microfossil groups and diatom species. The occurrence of siliceous microfossils is restricted to the uppermost few meters of the pelagic cap sediments. As a result of dissolution and silica diagenesis, the more dissolution-susceptible diatom valves and silicoflagellate skeletons are absent below 1–2.5 m core depth in the subtropical sites. In the highly calcareous and highly porous sediments on Limalok, Lo-En, and Wodejebato guyots, siliceous microfossils are now absent from lower Quaternary and upper Neogene sediments. Diatoms of slightly older age (lower Quaternary and Pliocene) are preserved on MIT Guyot. The inverse correlation of the abundance of authigenic zeolites with that of diatoms and other siliceous microfossils indicates that dissolution of biogenic silica provided silica for the formation of the zeolites.

Better preservation of diatoms was found farther north on Takuyo-Daisan Guyot just off Japan. The higher terrigenous component, together with dominant vesicular ash in the Quaternary to upper Pliocene sediments, as well as the higher primary productivity in this region, helped to preserve biogenic silica. Diatom abundance and preservation increased from south to north, which seems to be related to an increase of primary productivity in the west wind-influenced latitudes. Close sample spacing in the upper 70 cm of Site 880 allows the demonstration of changing eolian input through the last approximately 50,000 yr.

The sponge spicule types recovered from the guyots in sediments deposited in water depths between 1200 and 1500 m are also documented. The paleogeographic distribution of the diatom species occurring in the pelagic sediments are briefly discussed, but preservation was relatively poor.

Diatom biostratigraphy within the upper 7 m at Site 880 allows identification of the *Neodenticula seminae* and the *Rhizosolenia curvirostris* Zones of the North Pacific Diatom Zonation.

INTRODUCTION

During Ocean Drilling Program (ODP) Leg 144, Neogene sediments capping shallow-water limestones and/or volcanic rocks on five guyots in the western North Pacific (Table 1; see site map preceding title page) were cored. These drill holes provide a south-north transect from 5°N to 27°N and thus offer the chance to study the occurrence and geographic distribution of siliceous microfossils through the upper Neogene section over this latitudinal range.

MATERIAL

At the three southern sites between 5°N and 12°N, the upper Neogene sediments are yellowish gray to pinkish white, nannofossil-rich, foraminifer oozes, containing specks of manganese oxides and hydroxides in some cores. At these three sites, planktonic foraminifers are highly conspicuous components of the sediment, giving it its sandy texture. This, together with the high water content (80%–150%) of the sediments (Premoli Silva, Haggerty, Rack, et al., 1993) resulted in the liquefaction of the pelagic cap sediments from Holes 871A and 872A during handling of the cores on deck. The sediment-filled plastic liners were not completely filled, and parts of the sediment flowed back and forth within the liners. Thus, these cores cannot be used for detailed micropaleontological studies. When the cores for Holes 872C and 873B reached the deck, they were left in an upright position for several hours. In addition, a perforated lid (a "piglet") was placed on the sediment surface. Before the cores were split lengthwise, the water that had collected at the top was drained. This procedure hindered sediment mixing within the core liner and helped

to hold the sediment in place. These cores can be used for detailed micropaleontological analysis.

Farther north, at Site 878, pelagic sediments are rich in manganese micronodules and are yellowish brown in color. The water content of these sediments was only 50%–80%, much less than at the southern sites. At Site 880, the northernmost site drilled, the sediments were grayish brown to olive gray and contained a high percentage of terrigenous matter and volcaniclastic material. The sediments also had a much firmer consistency, having a water content between 46% and 70% (Premoli Silva, Haggerty, Rack, et al., 1993), so the use of the piglet was unnecessary.

METHODS

Between 10 and 25 g of sediment sample was necessary to obtain a sufficient number of siliceous microfossils of the Neogene foraminifer sands from Holes 871A, 872A, 872C, and 873B. Only 3–10 g was necessary at Site 878, and about 0.5 g was sufficient at Site 880. The sediment samples were prepared quantitatively so that the number of diatom valves per gram sediment could be determined. In addition, the weight of the total CH₃COOH-insoluble residue >2 µm was determined.

The weighed, dry samples were treated only with 10% CH₃COOH. Slides were prepared from the insoluble residue with a known aliquot of the sample and a statistically random distribution following the method of Battarbee (1973).

Hyrax was used as the embedding resin (refractive index $n = 1.71$; solvent: toluene). The slides were scanned at a 1000 \times magnification under a Zeiss Axioplan microscope using a 100 \times Plan-APOCHROMAT oil immersion objective with a numerical aperture of 1.40. If possible, a minimum of 300 siliceous microfossils, or 100 diatoms in Hole 880A, was counted per sample. The counts are related to the dry weight of the sediment and corrected for the salt content from the evaporated pore water.

¹ Haggerty, J.A., Premoli Silva, I., Rack, F., and McNutt, M.K. (Eds.), 1995. *Proc. ODP, Sci. Results*, 144: College Station, TX (Ocean Drilling Program).

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Table 1. Geographic location of Leg 144 holes studied and thickness of Neogene pelagic sediment column recovered.

Hole	Guyot	Latitude (N)	Longitude (E)	Water depth (m)	Recovered Neogene section (m)
871A	Limalok (Harrie)	05°33.432'	172°20.658'	1254.6	133.7
872A	Lo-En	10°05.850'	162°51.960'	1083.6	120.0
872C	Lo-En	10°05.620'	162°52.002'	1082.1	126.0
873B	Wodejebato	11°53.838'	164°55.230'	1334.0	53.0
878A	MIT	27°19.143'	151°53.028'	1323.2	3.2
880A	Takuyo-Daisan (Seiko)	34°12.053'	144°18.074'	1524.8	18.4

The microscopic identification of zeolites was checked in selected samples using X-ray diffraction analysis (H. Rösch, BGR, Hannover). The surface sediment sample from Hole 872A did not contain actual surface sediment. It consisted of older material that must have been washed out during handling of the core on deck and then settled on top. The results from this sample are included in the tables but were omitted from the figures.

RESULTS AND DISCUSSION

Siliceous Microfossil Groups

The main marine siliceous microfossil groups (diatoms, radiolarians, sponge spicules, and archaeomonadaceae) are present in the surface sediments at all sites. The most dissolution-susceptible silicoflagellates occur only sporadically. Pollen (tricolpate and bisaccate types) and dinoflagellates occasionally occur in the CH₃COOH-insoluble residue also (Pl. 2, Figs. 12–14).

Because the occurrence of siliceous microfossils is restricted to the surface near part of the pelagic cap sequence, only the upper few meters of each hole were analyzed for this study. The only exception to this is the northernmost site (Site 880).

At the sites in the tropical-subtropical belt (Figs. 1–3; Tables 2–5, 7–10), the abundance of all siliceous microfossils decreases significantly within the upper 1–3 m of the sediment column. Further downcore, the CH₃COOH-insoluble residue of the pelagic cap sediments is dominated by authigenic silicates. As in the sediments mentioned previously, quartz grains, volcanic ash, phosphatic remains, and Fe and Mn oxides and hydroxides also occur. The main zeolites are clinoptilolite and phillipsite. In these sediments, which are practically barren of siliceous microfossils, only sporadic archaeomonadaceae (the most dissolution-resistant siliceous microfossil group) and very rare single fragments of radiolarians or diatoms occur.

Overall, the abundance of the siliceous microfossil groups in the highly calcareous sediments at the three southern sites (Sites 871, 872, and 873 = 93%–98% CaCO₃) is very low. Even the highest abundance of diatom valves is only on the order of 10³ per gram of sediment. That of the silicoflagellates is 1 order of magnitude lower still. The abundance of diatoms at these sites drops below the recognizable limit (below 10^{1.5} valves per gram of sediment) within the uppermost 1.5–2 m of cored sediment. The more dissolution-resistant siliceous microfossil groups of radiolarians, sponge spicules, and Archaeomonadaceae can reach higher abundances (up to 10⁴ specimens per gram sediment) and can also be traced farther downcore.

Further evidence that the abundance of siliceous microfossils at these sites is controlled by dissolution is the downcore increase in fragmentation, etching, and partial recrystallization of these microfossils. The inverse correlation of the abundance of the siliceous microfossils and zeolites suggests that these authigenic silicates have drawn silica for their formation from the dissolution of biogenic opal.

At the three southern sites, the CH₃COOH-insoluble residue >2 µm, which is composed of wind-transported material (mainly volcanic ash) as well as biosiliceous particles, zeolites, and Fe and Mn oxides and hydroxides, is <1% of the bulk sediment. Further north at Site 878, the CH₃COOH-insoluble residue >2 µm is >10% (Table 10;

Fig. 4). According to the age determined for the pelagic cap sediments at this site by Erba (Premoli Silva, Haggerty, Rack, et al., 1993), Holocene and upper Pleistocene sediments are missing and lower Pleistocene and older sediments were recovered in Core 144-878A-1R. Volcanic ash particles are the dominant component (vesicular ash is rare) in the CH₃COOH-insoluble residue together with Fe and Mn oxides and hydroxides. The abundance of siliceous microfossils is still relatively low (on the order of 10³), and they persist into much older sediments than at the sites farther south.

At the northernmost site (Site 880 off Japan, on Takuyo Daisan Guyot), the situation is completely different. Here, the CH₃COOH-insoluble residue >2 µm makes up between 40% and 90% of the bulk sediment. Most of this residue is silt- to sand-sized vesicular ash, giving the sediment its gray color and causing the dark specks. Only two cores were recovered at this site. Siliceous microfossils are present here in much higher abundance than they are at the sites studied farther south (e.g., silicoflagellates reach abundances of 10⁵ skeletons per gram sediment). They occur down to the base of the cored interval (Fig. 5). This, combined with the absence of zeolites, indicates that silica diagenesis was much slower than at the more southern sites. Nevertheless, a distinct stepwise decrease in the preservation of biosiliceous remains is also recognizable at this site. A decrease in diatom abundance (from an average of 10⁵ to <10⁴ valves/g sediment) occurs at 7 mbsf in the core within the *Pseudoemiliania lacunosa* Zone (Erba, this volume). Also, the abundance of silicoflagellates and radiolarians decreases at this core depth (Fig. 5). The occurrence of *Spermatogonia* sp. and *Actiniscus* sp. (dinoflagellates with internal siliceous spicules) is restricted to these upper 7 m of the sediment column. Here, diatom abundance fluctuates strongly between 10³ and 10⁶ valves per gram sediment. Below 7 m core depth, only the most dissolution-resistant diatom valves or fragments and the more robust radiolarian species are left. Variations in diatom abundance are comparatively small (between 10³ and 10⁴ valves/g sediment). A second step in the decrease of preservation is reflected at 15 m core depth by the relative increase of the more dissolution-resistant radiolarians and freshwater diatoms as compared with the less dissolution-resistant marine diatoms (Fig. 6).

The overall much better preservation of siliceous microfossils at Site 880 may result in part from a higher productivity in these waters just off Japan. On the other hand, dissolved silica from the abundant vesicular ash in these sediments may help preserve the biogenic silica. Unfortunately, no analyses of dissolved silica in the pore waters of Hole 880A are available. For the sandy foraminifer oozes on Limalok, Lo-En, and Wodejebato guyots, pore-water analyses yielded characteristic seawater values and no downcore gradients. This indicates that pore-water flow, and possibly flushing of the sediment column during core recovery, controlled the concentration of ions in dissolution (Opdyke, this volume), but diffusion did not.

A close sample spacing was obtained of one sample every 2–3 cm for the upper 66 cm of core, allowing for checking of climate-related changes in the abundance and preservation of siliceous microfossils.

A special characteristic of Site 880 is the relatively high and frequent occurrence of phytoliths and freshwater diatoms (Fig. 5; Table 6; Pl. 2, Figs. 15–25), including members of the *Aulacoseira granulata* group, *Stephanodiscus* sp., and a single specimen each of *Navicula* sp., *Fragilaria* sp., *Cocconeis placentula*, *Cymbella* sp., *Cyclotella* sp., *Eunotia* sp., *Epithemia* sp., and *Achnanthes* sp. Pollen grains were also frequently found in samples from this site (Table 6). These three component groups suggest a definite input of windblown material by west winds. The upper 70 cm of this hole belongs to the *Emiliania huxleyi* Acme Zone (Erba, this volume); thus, this segment is younger than 70,000 yr. The abundance changes of vesicular ash and the different siliceous microfossil groups (except for freshwater diatoms) all show a similar trend (Figs. 5, 7). Their abundances drop slightly and gradually upsection, reaching minimum values at a core depth of 20 to 5 cm. In the upper 5 cm, the abundances of all these components strongly increase, in most cases by more than 1 order of magnitude.

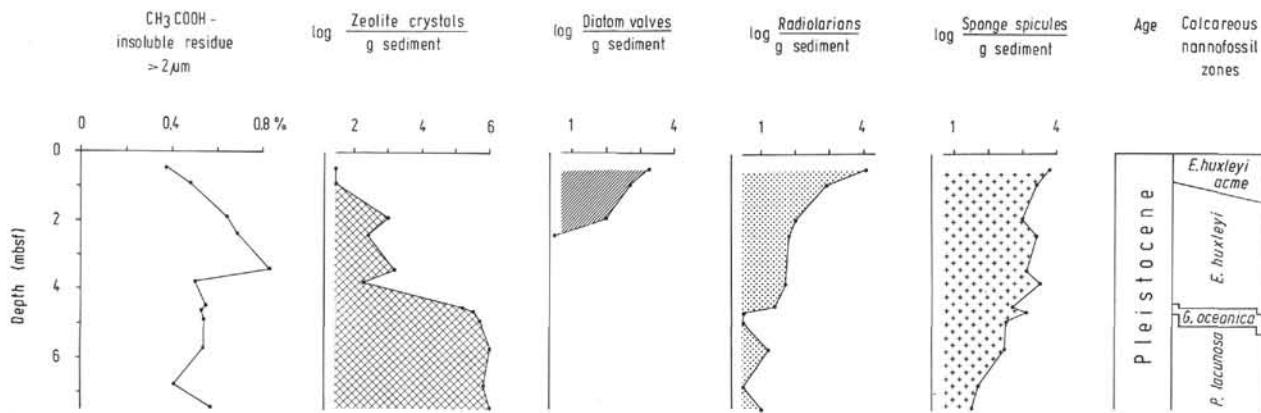


Figure 1. Abundance of the CH_3COOH -insoluble residue $>2 \mu\text{m}$, zeolites, and major siliceous microfossil groups in upper Quaternary sediments from Hole 871A. Calcareous nannofossil zones after Premoli Silva, Haggerty, Rack, et al. (1993).

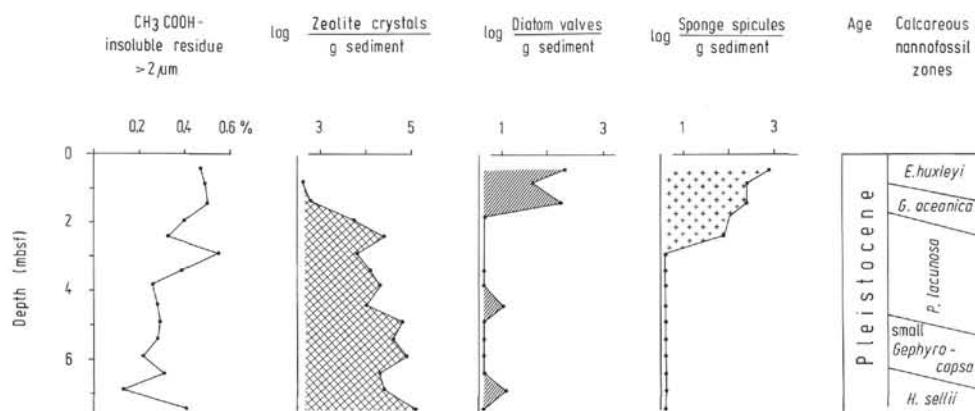


Figure 2. Abundance of the CH_3COOH -insoluble residue $>2 \mu\text{m}$, zeolites, and major siliceous microfossil groups in Quaternary sediments from Hole 872A. Calcareous nannofossil zones after Premoli Silva, Haggerty, Rack, et al. (1993).

Table 2. Occurrence of diatom species in samples from the upper Quaternary pelagic sediments, Hole 871A.

Core, section, interval (cm)	Siliceous sponge spicules	Radiolarians	Archaeomonadaceae	Silicoflagellates	Dinoflagellates	Fillings of worm burrows	Pollen	Asterolampra marylandica	Asteromphalus sacrophagus forma obovatus	Asteptia nodulifer	Asteptia tabularis	Chaeoceros messanensis	Cyclotella littoralis	Ethmodiscus sp. fragments	Hemidiscus cuneiformis	Nitzschia interruptostriata	Nitzschia marina	Rhizosolenia heronii	Rhizosolenia hebetata	Thalassionema nitzschiae	Thalassiothrix longissima
144-871A-																					
IH-1, 41-43	C	D	R	S	S	R	R	—	R	C	F	R	—	R	F	—	R	—	R	F	
IH-1, 87-91	D	C	R	—	S	R	R	—	—	R	—	—	—	—	—	—	—	—	—	—	
IH-2, 41-43	C	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-2, 87-91	C	S	—	—	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-3, 41-43	C	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-3, 74-78	C	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-3, 148-150	R	S	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-4, 14-18	F	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-4, 39-41	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-4, 121-123	R	S	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-5, 75-79	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
IH-CC	S	S	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	

Notes: The abundance of each species and of the microfossil groups is given as follows: S (scarce) = <1%, R (rare) = 1%-3%, F (few) = 3%-10%, C (common) = 10%-30%, and D (dominant) = >30%.

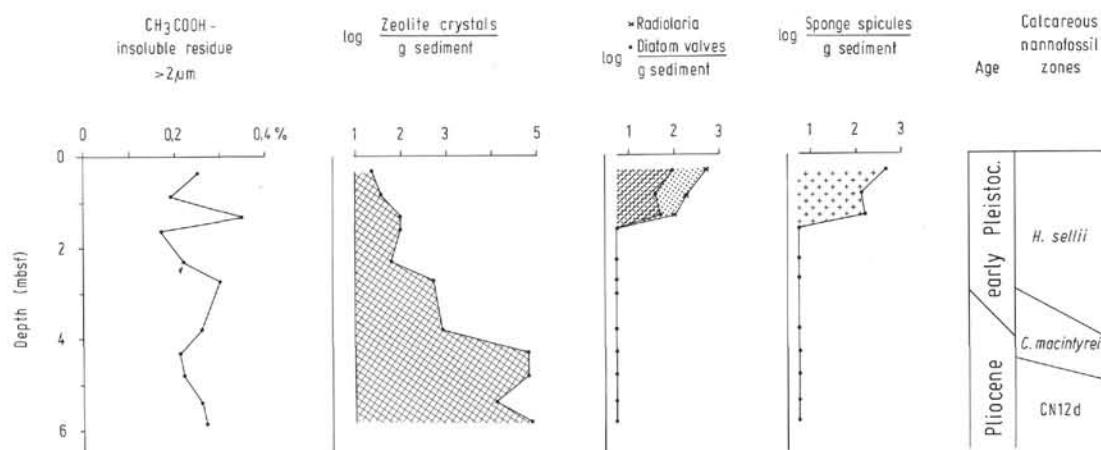


Figure 3. Abundance of the CH_3COOH -insoluble residue $>2 \mu\text{m}$, zeolites, and major siliceous microfossil groups in Quaternary sediments from Hole 873B. Calcareous nannofossil zones after Premoli Silva, Haggerty, Rack, et al. (1993).

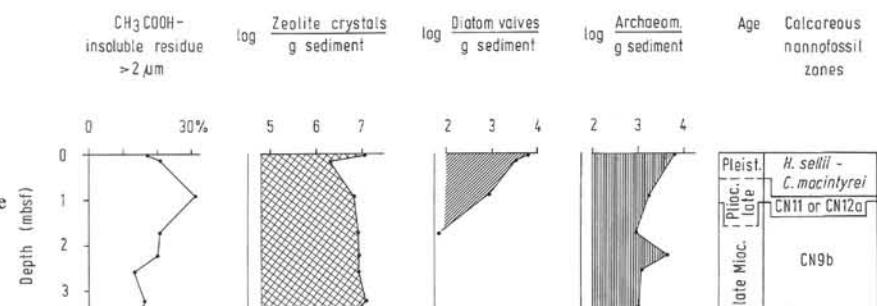


Figure 4. Abundance of the CH_3COOH -insoluble residue $>2 \mu\text{m}$, zeolites, and major siliceous microfossil groups in Neogene sediments from Hole 878A. Calcareous nannofossil zones after Premoli Silva, Haggerty, Rack, et al. (1993).

Table 3. Occurrence of diatom species in samples from the Quaternary pelagic sediments, Holes 872A and 872C.

Core, section, interval (cm)	Archaeomonadaceae	Siliceous sponge spicules	Radiolarians	Silicoflagellates	Dinoflagellates	Pollen	<i>Asteromphalus hookeri</i>	<i>Apeititia nodulifer</i>	<i>Apeititia tabularis</i>	<i>Ehmdiscus</i> sp. fragments	<i>Hemidiscus cuneiformis</i>	<i>Neodenticula seminae</i>	<i>Nitzschia bicapitata</i>	<i>Nitzschia marina</i>	<i>Rhizosolenia bergonii</i>	<i>Thalassionema nitzschioides</i>	<i>T. nitzschioides</i> var. <i>parva</i>	<i>Thalassiosira</i> sp.	<i>Thalassiothrix longissima</i>	Age	Calcareous nannofossil zones
144-872A-																					
1H-1 (top)	S	S	R	C	—	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-1, 39-43	F	C	F	F	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-1, 89-93	R	F	R	F	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-1, 140-144	R	R	R	F	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-2, 39-43	R	R	R	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-2, 89-93	R	R	R	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-2, 140-144	—	—	—	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-3, 39-43	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-3, 89-93	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-3, 140-144	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-4, 39-43	R	—	—	—	—	—	S	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-4, 89-93	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
1H-4, 140-144	R	—	—	—	—	—	R	S	—	—	—	—	—	—	—	—	—	—	—	—	
1H-5, 39-43	R	—	—	—	—	—	R	R	—	—	—	—	—	—	—	—	—	—	—	—	
1H-5, 89-93	R	—	—	—	—	—	R	R	—	—	—	—	—	—	—	—	—	—	—	—	
1H-CC	C	F	C	—	R	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	
144-872C-																					
1H-1, 0-10	F	C	C	—	—	—	—	S	—	—	—	—	—	—	—	—	—	S	S	—	
1H-1, 16-20	R	C	C	—	—	—	S	S	S	S	—	—	—	—	S	—	—	—	—	—	
1H-1, 36-42	C	F	C	—	—	—	S	S	S	S	R	—	—	S	—	—	—	—	—	—	
1H-1, 52-56	F	F	C	—	—	—	S	—	—	—	S	—	—	S	—	—	—	—	—	—	
1H-1, 75-79	F	F	F	—	—	—	S	—	—	—	S	—	—	S	—	—	—	—	—	—	
1H-1, 90-94	C	F	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	

Note: See Table 2 for explanation of abbreviations.

Table 4. Occurrence of diatom species in samples from the Quaternary pelagic sediments, Hole 873B.

Core, section, interval (cm)	Siliceous sponge spicules	Archaeomonadaceae	Siliceous sponge spicules	Radiolarians	Silicoflagellates	Pollen	<i>Actinocyclus</i>	<i>Actinocyclus curvatus</i>	<i>Asteromphalus robustus</i>	<i>Asteromphalus sacrophagus forma sacrophagus</i>	<i>Appeitia africana</i>	<i>Appeitia nodulifer</i>	<i>Cyclotella</i>	<i>Ethmodiscus</i> sp. fragments	<i>Hemidiscus</i> sp.	<i>Liosphaerionia</i> of <i>Appeitia</i> sp.	<i>Nitzschia marina</i>	<i>Nitzschia bicapitata</i>	<i>Paralia sulcata</i>	<i>Roperia tessellata</i>	<i>Thalassionema bacillaris</i>	<i>T. nitzschioides</i> var. <i>parva</i>	<i>Thalassiosira eccentrica</i>	<i>Thalassiosira leptopus</i>	<i>Thalassiosira oestrupii</i>	<i>Thalassiosira</i> sp. 1	
144-873B-																											
1H-1, 0-2		R		F	A	S	R																				
1H-1, 30-34		R		C	C	—	R																				
1H-1, 81-86				C	C	—																					
1H-2, 0-2				C	C	S	—																				
1H-2, 30-34				—	—	—	R																				
1H-2, 101-106																											
1H-3, 21-25																											
1H-3, 130-134																											
1H-4, 32-36																											
1H-4, 82-86																											
1H-5, 9-14																											
1H-CC		R	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	

Notes: See Table 2 for explanation of abbreviations. Single asterisk (*) = only Mn-hydroxide crusts from former microfossils are left.

Table 5. Occurrence of diatom species in samples from the Neogene pelagic sediments, Hole 878A.

Core, section, interval (cm)	Siliceous sponge spicules	Radiolarians	Archaeomonadaceae	Silicoflagellates	Fillings of worm burrows	Pollen	<i>Actinocyclus ehrenbergii</i>	<i>Asterolampra marylandica</i>	<i>Asteromphalus hiltonianus</i>	<i>Asteromphalus ornithopas</i>	<i>Appeitia africana</i>	<i>Appeitia neocrenulata</i>	<i>Appeitia nodulifer</i>	<i>Chaetoceros messanensis</i>	<i>Coscinodiscus argus</i>	<i>Coscinodiscus asteromphalus</i>	<i>Coscinodiscus pacificus</i>	<i>Coscinodiscus radiatus</i>	<i>Ethmodiscus</i> sp. fragments	<i>Fragilariopsis doliolus</i>	<i>Hemidiscus cuneiformis</i>	<i>Nitzschia interruptostriata</i>	<i>Nitzschia marina</i>	<i>Paralia sulcata</i>	<i>Planktoniella sol</i>	<i>Pseudoriveratum</i> sp.	<i>Rhizosolenia bergonii</i>	<i>Rhizosolenia hebetata</i>	<i>Roperia tessellata</i>	<i>Thalassionema bacillaris</i>	<i>T. nitzschioides</i> var. <i>parva</i>	<i>Thalassiosira eccentrica</i>	<i>Thalassiosira leptopus</i>	<i>Thalassiosira oestrupii</i>	<i>Thalassiosira</i> sp.	<i>Thalassiotrix longissima</i>
144-878A-																																				
IR-1 (top)	R	A	F	R	R	—	R	S	S	S	R	S	F	R	S	S	S	S	S	S	R	S	S	S	S	S	S	S	S	S						
IR-1, 12-14	R	—	S	—	R	—	—	—	—	—	—	S	S	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							
IR-1, 93-95	—	—	R	—	R	—	—	—	—	—	—	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							
IR-2, 21-23	—	R	R	—	R	—	—	—	—	—	—	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							
IR-2, 71-73	S	S	R	—	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							
IR-2, 108-110	—	—	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							
IR-3, 21-23	—	—	—	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							
IR-3, 35-37	—	—	—	R	S	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							
IR-CC	—	—	—	R	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—	—							

Note: See Table 2 for explanation of abbreviations.

Because of the covariance of components from different sources (e.g., marine planktonic diatoms and vesicular ash), the abundance changes of the siliceous microfossil groups are not controlled by dilution with volcanic ash (Figs. 5, 6). Rather, the more ash present, the greater the number of siliceous microfossils, the higher their diversity, and the better their preservation. The abundance changes may result from two factors: (1) the abundant presence of vesicular ash in the sediments helps preserve siliceous microfossils and (2) the abundance fluctuations of volcanic ash and biosiliceous particles may be controlled by the same cause. This cause could be climate-related changes in the west wind intensity or in aridity and, consequently, in dilution by windblown detritus. However, it could also be dilution by CaCO_3 , thus changing the productivity of biogenic carbonate producers. Unfortunately, no detailed CaCO_3 determinations yet exist for this hole. The ratio of freshwater to marine diatoms (Fig. 7)

indicates that windblown detritus originating from west of the site has increased during the time interval represented in the core between 20 and 5 cm depth bsf.

Isotopic studies would have been helpful for the interpretation. In their absence, this interval is interpreted to represent the last glacial maximum and Isotopic Stage 2. The rise in abundance in the upper 5 cm is interpreted to represent Termination I, and most of Isotopic Stage 1 would be missing. If this interpretation is correct, the time interval of the last glacial maximum was a period of increased wind intensity and of increased aridity in the coastal mainland west of the site, thus forming a source area for the windblown dust deposited offshore.

Below 66 cm, the spacing of the samples is too far apart to allow such study. Here, a clear cyclicity in abundances of eolian components that would indicate changes between more arid and more humid conditions in the hinterland cannot be recognized. Abundance of

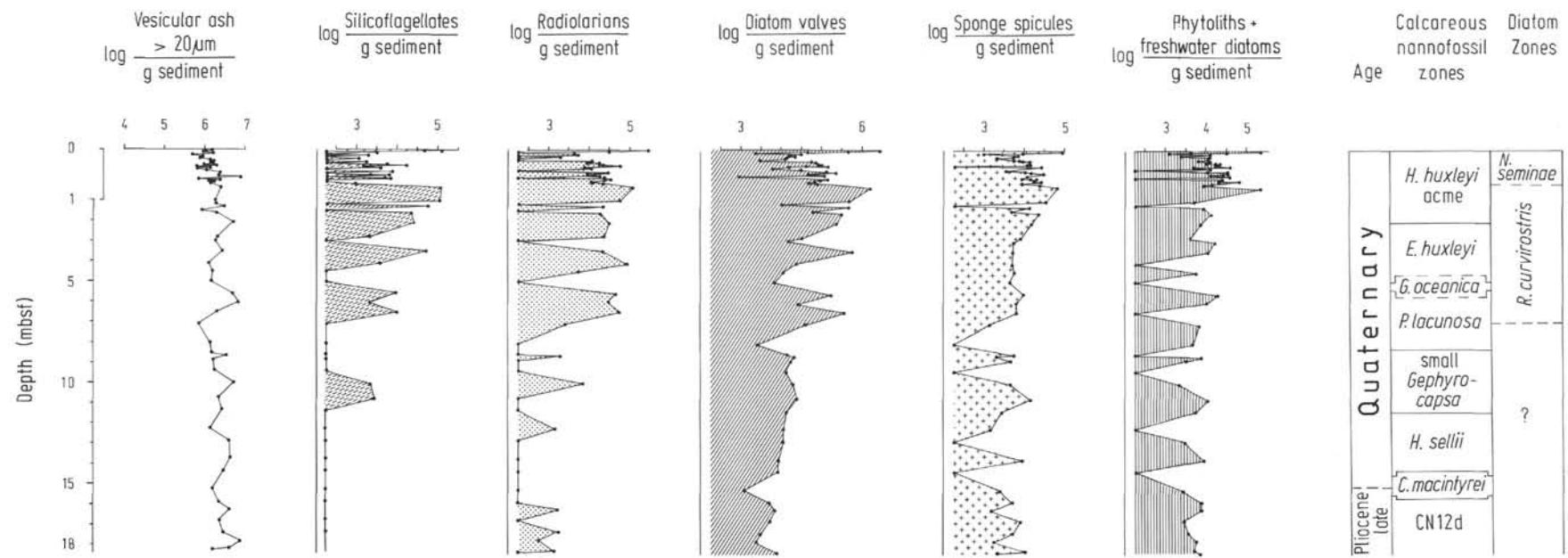


Figure 5. Abundance of vesicular ash >20 µm and the major siliceous microfossil groups in upper Pliocene to Quaternary sediments from Hole 880A plotted on a logarithmic scale. Note the different scale for the uppermost meter of the core depth. Calcareous nannofossil zones after Premoli Silva, Haggerty, Rack, et al. (1993).

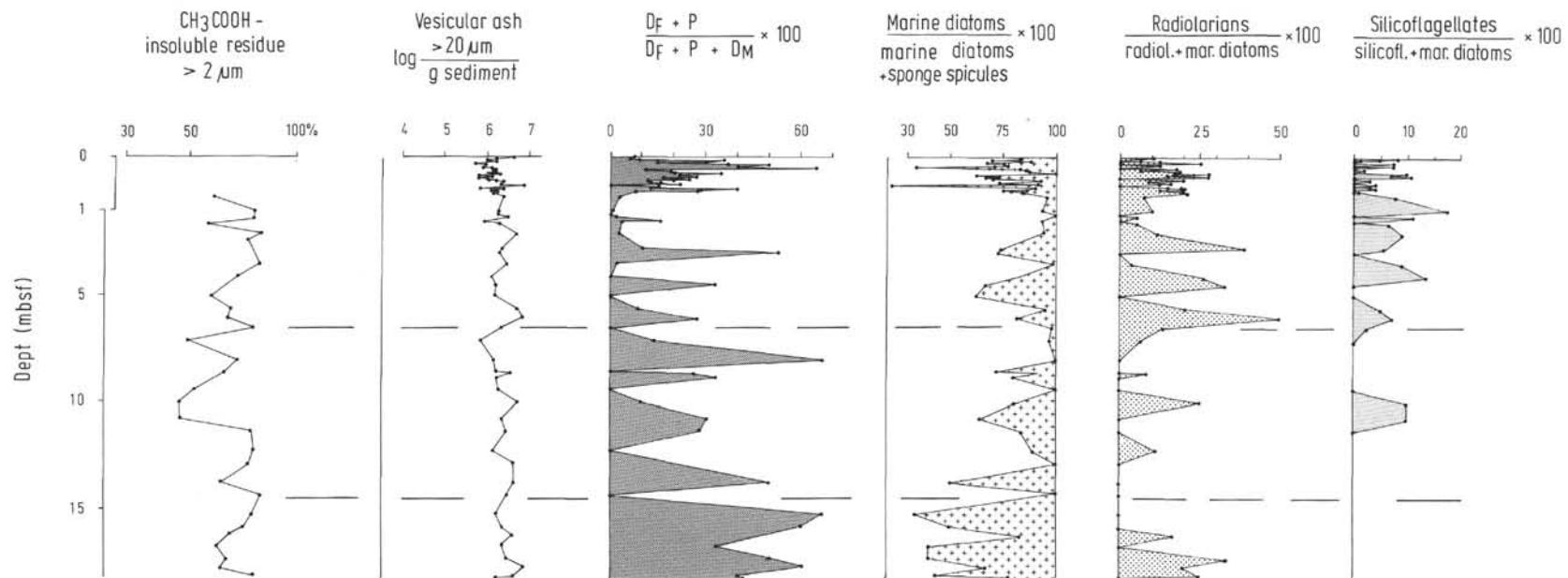


Figure 6. Abundance of vesicular ash $> 20 \mu\text{m}$, CH_3COOH -insoluble residue $> 2 \mu\text{m}$, and siliceous microfossil groups of different dissolution-susceptibility and of nonmarine vs. marine diatoms (D_M = marine diatoms, D_F = freshwater diatoms, P = phytoliths) in Hole 880A. Note the different scale for the core depth in the uppermost meter. The dashed lines mark steps of decreasing preservation of biogenic silica.

Table 6. Occurrence of diatom species in samples from the upper Pliocene and Quaternary pelagic sediments, Hole 880A.

Table 6 (continued).

	Core, section, interval (cm)			
1H-CC, upper part	S	Siliceous sponge spicules		
1H-CC, lower part	R	—	Radiolarians	
2H-1, 48–49	R	S	Archaeomonadaceae	
2H-1, 108–109	R	R	Silicoflagellates	
2H-2, 49–50	R	R	<i>Actiniscus</i> sp.	
2H-2, 109–110	S	S	<i>Spermatogonia</i> sp.	
2H-3, 48–49	S	S	Pollen	
2H-3, 109–110	R	S	Fish scale	
2H-4, 49–50	R	—	<i>Actinoecytus curvatus</i>	
2H-4, 107–108	S	—	<i>A. ellipticus</i>	
2H-5, 49–50	R	—	<i>A. ingens</i>	
2H-5, 108–109	R	R	<i>A. octonarius</i>	
2H-5, 140–150	R	S	<i>A. tenellus</i>	
2H-6, 49–50	R	S	<i>A. sp. 1</i>	
2H-6, 109–110	R	S	<i>Actinopychus senarius</i>	
2H-6, 142–150	S	S	<i>A. splendens</i>	
2H-7, 49–50	R	S	<i>Asterolampria marylandica</i>	
2H-CC	R	F	<i>Asteromphalus elegans</i>	
			<i>A. hookeri</i>	
			<i>Arpeitia africana</i>	
			<i>A. neocrenulata</i>	
			<i>A. nodulifer</i>	
			<i>A. tabularis</i>	
			<i>Bacteriastrom sp.</i>	
			<i>Chaetoceros messanensis</i>	
			<i>Chaetoceros sp.</i>	
			<i>Coscinodiscus argus</i>	
			<i>C. gigas</i>	
			<i>C. marginatus</i>	
			<i>C. pacificus</i>	
			<i>C. radiatus</i>	
			<i>C. sp.</i>	
			<i>Cyclotella littoralis</i>	
			<i>Denticulopsis lauta</i>	
			<i>Diploneis bombus</i>	
			<i>Ethmodiscus</i> sp. fragments	
			<i>Fragilariaopsis dolichos</i>	
			<i>Hemidiscus caniformis</i>	
			<i>H. caniformis</i> var.	
			<i>Hemidiscus</i> sp.	
			<i>Navicula directa</i>	
			<i>Neodenticula kamtschatka</i>	
			<i>N. seminae</i>	
			<i>Nitzschia bicapitata</i>	
			<i>N. braunii</i>	
			<i>N. interrupta</i>	
			<i>N. kolaczekii</i>	
			<i>N. marina</i>	
			<i>N. sicula</i>	
			<i>N. sp.</i>	

Table 6 (continued).

Core, section, interval (cm)	Freshwater diatoms																													
	Freshwater diatoms																													
	<i>Odonella aurita</i>																													
	<i>Paralia sulcata</i>																													
	<i>Planktoniella sol</i>																													
	<i>Pseudodictic. cinnamomeum</i>																													
	<i>Pyxilla reticulata</i>																													
	<i>Rhaphoneis surireloides</i>																													
	<i>Rhizosolenia bergonii</i>																													
	<i>Rhizosolenia hebetaria</i>																													
	<i>Rhizosolenia setigera</i>																													
	<i>Rhizosolenia styliformis</i>																													
	<i>Rhizosolenia styliformis</i> var. <i>longispina</i>																													
	<i>Ropeira tessellata</i>																													
	<i>Rosselia tatsunokuchiensis</i>																													
	<i>Rouxia californica</i>																													
	<i>Simonsenella barbata</i>																													
	<i>Stellaria microtrias</i>																													
	<i>Stephanopyxis turris</i>																													
	<i>Thalassionema bacillaris</i>																													
	<i>Thalassionema nitescioides</i>																													
	<i>T. nitescioides</i> var. <i>vobusa</i>																													
	<i>T. nitescioides</i> var. <i>parva</i>																													
	<i>Thalassiosira angustelineata</i>																													
	<i>Thalassiosira cf. decipiens</i>																													
	<i>Thalassiosira eccentrica</i>																													
	<i>Thalassiosira ferelinata</i>																													
	<i>Thalassiosira gravida</i>																													
	<i>T. joussea (T. nidulus)</i>																													
	<i>T. latimarginata (T. irifluita)</i>																													
	<i>Thalassiosira leptopus</i>																													
	<i>T. lineata</i>																													
	<i>T. aeruginosa var. oestrupii</i>																													
	<i>T. aeruginosa var. plana</i>																													
	<i>T. sp. 1</i>																													
	<i>T. sp.</i>																													
	<i>Thalassiosira longissima</i>																													
	<i>Xanthiopavis</i> sp.																													
	Other resting spore types																													
	<i>Achnanthes</i> sp.																													
	<i>Aulacoseira</i> sp.																													
	<i>Cocconeis</i> sp.																													
	<i>Cyclotella</i> radioosa																													
	<i>Cymbella</i> sp.																													
	<i>Diploneis</i> sp.																													
	<i>Epithemia</i> sp.																													
	<i>Eunotia</i> sp.																													
	<i>Fragilaria</i> sp.																													
	<i>Navicula</i> sp.																													
	<i>Stephanodiscus</i> sp.																													

Table 6 (continued).

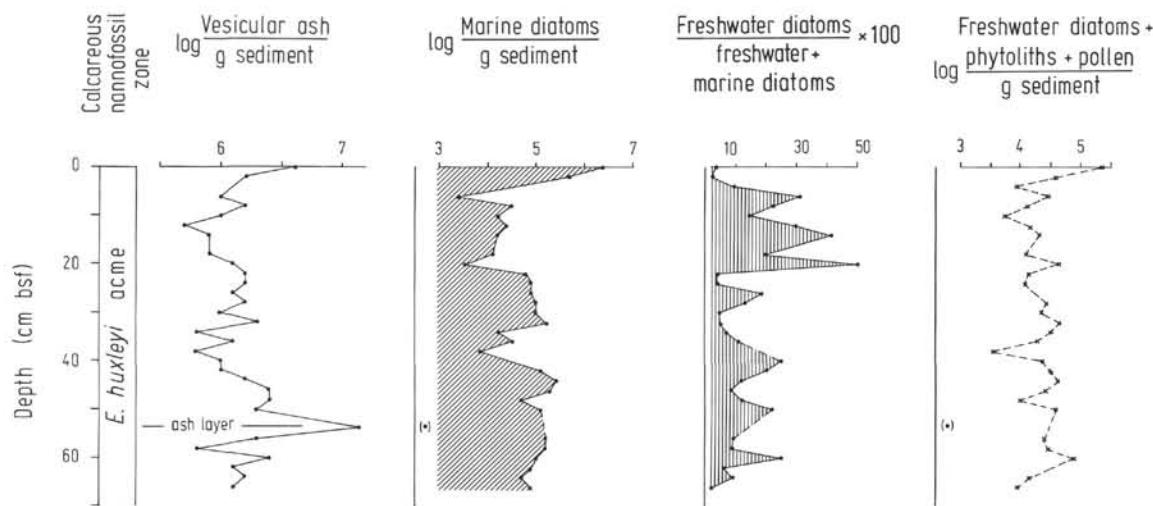


Figure 7. Enlarged plot for the upper 70 cm of Hole 880A showing the abundance of vesicular ash $>20\mu\text{m}$, marine diatoms, and eolian transported freshwater diatoms, phytoliths, and pollen. All are plotted on a logarithmic scale versus the ratio of freshwater to marine diatoms.

Table 7. Abundance of major siliceous microfossil groups, zeolites, and CH_3COOH -insoluble residue $>2\mu\text{m}$ in samples from Hole 871A.

Core, section, interval (cm)	Depth (cm bsf)	CH_3COOH -insoluble residue (%)	Archaeomonadaceae	Siliceous sponge spicules	Radiolarians	Silicoflagellates	Diatoms	Log zeolite crystals per gram sediment
144-871A-								
IH-1, 41-43	41-43	0.4	2.8	3.8	4.1	2.0	3.3	—
IH-1, 87-91	87-91	0.5	2.9	3.4	2.9	—	2.7	—
IH-2, 41-43	191-193	0.6	—	3.0	2.0	—	2.0	3.0
IH-2, 87-91	237-241	0.7	—	3.4	1.8	—	—	2.4
IH-3, 41-43	341-343	0.8	—	3.1	1.7	—	—	3.2
IH-3, 74-78	374-378	0.5	—	3.5	1.7	—	—	2.3
IH-3, 148-150	448-450	0.6	2.9	2.7	1.4	—	—	5.2
IH-4, 14-18	464-468	0.5	—	3.1	—	—	—	5.5
IH-4, 39-41	489-491	0.5	—	2.5	—	—	—	5.7
IH-4, 121-123	571-573	0.5	1.9	2.5	1.2	—	—	6.0
IH-5, 75-79	675-679	0.4	—	1.7	—	—	—	5.8
IH-CC	745	0.6	1.9	1.5	1.0	—	—	6.0

Note: Major siliceous microfossil group values calculated as log siliceous microfossil group per gram sediment.

Table 8. Abundance of major siliceous microfossil groups, zeolites, and CH_3COOH -insoluble residue $>2\mu\text{m}$ in samples from Holes 872A and 872C.

Core, section, interval (cm)	Depth (cm bsf)	CH_3COOH -insoluble residue (%)	Archaeomonadaceae	Siliceous sponge spicules	Radiolarians	Silicoflagellates	Diatoms	Log zeolite crystals per gram sediment
144-872A-								
IH-1 (top)	0-4	0.3	—	—	—	—	—	4.8
IH-1, 39-43	39-43	0.5	2.3	2.9	2.9	—	2.4	—
IH-1, 89-93	89-93	0.5	2.0	2.4	2.4	—	1.7	2.6
IH-1, 140-144	140-144	0.5	2.4	2.4	1.8	—	2.3	2.8
IH-2, 39-43	189-193	0.4	2.0	2.0	2.4	—	—	3.7
IH-2, 89-93	239-243	0.3	1.6	1.9	—	—	—	4.4
IH-2, 140-144	290-294	0.6	—	—	1.2	—	—	3.8
IH-3, 39-43	339-343	0.4	2.1	—	—	—	—	4.1
IH-3, 89-93	389-393	0.3	2.1	—	—	—	—	4.3
IH-3, 140-144	440-444	0.3	—	—	—	—	1.0	4.0
IH-4, 39-43	489-493	0.3	2.0	—	—	—	—	4.8
IH-4, 89-93	539-543	0.3	2.0	—	—	—	—	4.6
IH-4, 140-144	590-594	0.2	1.8	—	—	—	—	4.9
IH-5, 39-43	639-643	0.3	2.3	—	—	—	—	4.3
IH-5, 89-93	689-693	0.1	1.9	—	—	—	1.1	4.4
IH-CC	745	0.4	2.4	—	—	—	—	5.1
144-872C-								
IH-1, 0-10	0-10	0.4	2.7	2.6	2.8	—	1.7	2.7
IH-1, 16-20	16-20	0.6	2.2	2.7	2.8	—	2.0	2.4
IH-1, 36-42	36-42	0.6	3.2	2.8	3.1	—	1.9	3.2
IH-1, 52-56	52-56	0.5	2.8	2.5	2.9	—	2.4	2.7
IH-1, 75-79	75-79	0.4	2.6	2.4	2.6	—	2.0	2.7
IH-1, 90-94	90-94	0.7	3.1	2.7	2.1	—	1.2	2.8

Note: Major siliceous microfossil group values calculated as log siliceous microfossil group per gram sediment.

Table 9. Abundance of major siliceous microfossil groups, zeolites, and CH_3COOH -insoluble residue >2 μm in samples from Hole 873B.

Core, section, interval (cm)	Depth (cm bsf)	CH_3COOH -insoluble residue (%)	Archaeomonadaceae	Siliceous sponge spicules	Radiolarians	Silicoflagellates	Diatoms	Log zeolite crystals per gram sediment
144-873B-								
IH-1, 30-34	30-34	0.25	1.4	2.7	—	—	—	2.0
IH-1, 81-86	81-86	0.19	—	2.2	2.3	—	—	1.6
IH-2, 0-2	130-132	0.35	—	2.2	2.0	—	—	2.0
IH-2, 30-34	160-164	0.17	1.2	—	—	—	—	2.0
IH-2, 101-106	231-236	0.22	—	—	—	—	—	1.8
IH-3, 21-25	271-275	0.30	—	—	—	—	—	2.7
IH-3, 130-134	380-384	0.26	—	—	—	—	—	2.9
IH-4, 32-36	432-436	0.21	—	—	—	—	—	4.8
IH-4, 82-86	482-486	0.22	—	—	—	—	—	4.8
IH-5, 9-14	539-544	0.26	—	—	—	—	—	4.1
IH-CC	585	0.27	—	—	—	—	—	4.9

Notes: Major siliceous microfossil group values calculated as log siliceous microfossil group per gram sediment. Single asterisk (*) = manganese coating.

freshwater diatoms shows strong fluctuations down to 15 m core depth (Figs. 5, 6), therefore, throughout the Quaternary represented in this hole.

Siliceous Sponge Spicules

Many different types of sponge spicules, microscleres, and macroscleres were found in the upper Quaternary sediments (Pls. 2 and 3). The nomenclature used here follows Wiedenmayer (1977). A documentation of the types of siliceous sponge spicules from sediments on top of guyots is of special interest because it provides information on which species of sponges live in this region at water depths between 1000 and 1500 m. On continental slopes, downslope transport from environments such as reefs can distort the depth distribution; however, on a freestanding guyot, the admixture of allochthonous material from shallower environments is not of concern. For the Eocene to the Pleistocene, Ivanik (1983) and McCartney (1990) documented sponge spicules from deep-sea sediments deposited in similar locations on seamounts or submarine plateaus (Maurice Ewing Bank, Maud Rise). Deep-sea sponge studies containing a description of spicules include Schulze (1899), Lévi (1964), and Tendal (1973). Oxea, strongyle, style, acanthostyloge, acanthooxea, acanthostyle, tylostyle, and microhabd were dominant among the macroscleres. The microsclere type sigma was also found relatively frequently as were isochele, spheraster, euaster, and fragments of fibroreticulate skeletons of hexactinellids. Tetraxons such as calthrops or oxyhexaster were very rare as were the microsclere types of hexactinellids (such as pinule and scropule).

Diatoms

Sites 871 and 878

On the three western Pacific guyots in the tropical-subtropical region between 5°N and 12°N, the diversity of diatoms is very low and the abundance is scarce (Tables 2-4). At Site 878 (27°19.14'N), diversity is slightly higher (Table 5). At Sites 871, 872, 873, and 878, the relatively dissolution-resistant and geographically widely distributed species *Thalassiothrix longissima*, *Azpeitia nodulifer*, *A. neocrenulata*, *Thalassionema nitzschiooides*, and *Nitzschia marina* are most commonly found, as well as fragments of *Ethmodiscus* sp., and sporadic *Rhizosolenia bergenii*, *Cyclotella litoralis*, *Hemidiscus cuneiformis*, and *Chaetoceros messanensis*. Some of these are illustrated on Plates 1 and 4.

Species that are predominantly tropical-subtropical, such as *Asteromphalus sacrophagus* forma *obovatus*, and also *A. elegans*, *A. hiltonianus*, and *Asterolampra marylandica*, occur only rarely.

Also, species that have a wider latitudinal range but which prefer more fertile waters, such as *Roperia tesselata*, *Nitzschia bicapitata*, *N. interruptestriata*, *Asteromphalus hookeri*, and *Thalassiosira* species, occur only sporadically at these western Pacific open-ocean locations.

No stratigraphic determination based on diatoms was possible at these sites because of the few species present and the absence of *Fragilariaopsis doliolus*, the name-giving species for the uppermost diatom zone for the low latitudes.

Site 880

On Takuyo-Daisan (Site 880), at approximately 34°N, the situation is quite different. This site lies within the area of increased planktonic primary productivity (see Koblentz-Mishke et al., 1970). Here, diatom abundance in the upper Quaternary sediments fluctuates between 10^4 to 10^6 valves per gram sediment (Figs. 5, 7; Table 11). The more dissolution-resistant valves such as *Xanthiopyxis* sp. and other resting spore types commonly dominate the assemblages. Valves of *Thalassiothrix longissima*, *Thalassionema nitzschiooides*, *T. nitzschiooides* var. *parva*, *Thalassiosira oestrupii*, *Cyclotella litoralis*, and *Chaetoc-*

Table 10. Abundance of major siliceous microfossil groups, zeolites, and CH₃COOH-insoluble residue >2 µm in samples from Hole 878A.

Core, section, interval (cm)	Depth (cm bsf)	CH ₃ COOH-insoluble residue (%)	Archaeomonadaceae	Siliceous sponge spicules	Radiolarians	Silicoflagellates	Diatoms	Log zeolite crystals per gram sediment
144-878A-								
IR (top)	0-1	17.78	3.8	—	3.1	3.1	3.8	7.0
IR-1, 12-14	12-14	21.79	2.5	3.3	—	—	3.5	6.3
IR-1, 93-95	93-95	32.03	3.3	—	—	—	2.9	6.8
IR-2, 21-23	171-173	21.10	4.0	—	—	—	—	6.9
IR-2, 71-73	221-223	20.08	3.7	2.3	—	—	—	6.9
IR-2, 108-110	258-260	14.14	3.1	—	—	—	—	6.9
IR-3, 21-23	321-323	16.61	—	—	—	—	—	7.1
IR-3, 35-37	335-337	16.31	3.0	—	—	—	—	7.0
IR-CC	350-352	21.26	3.1	—	—	—	—	7.0

Note: Major siliceous microfossil group values calculated as log siliceous microfossil group per gram sediment.

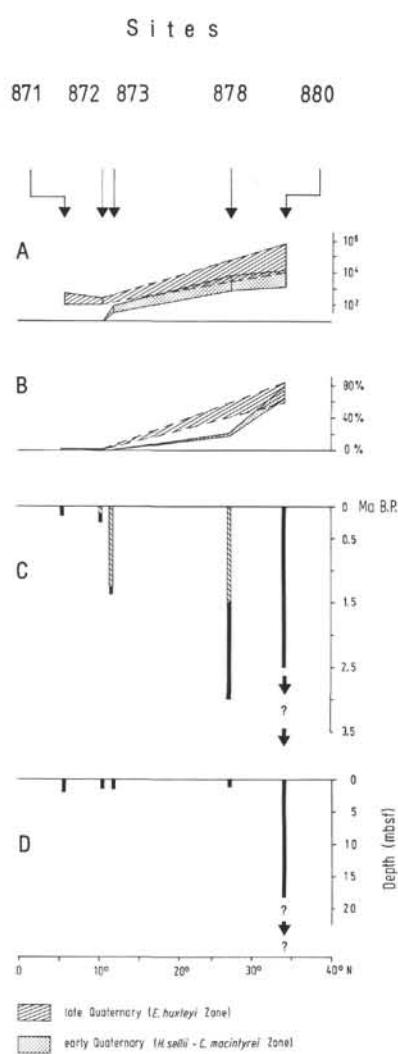


Figure 8. Latitudinal plot of selected parameters for two different time intervals in the Quaternary: the late Quaternary *Emiliania huxleyi* Zone, and the early Quaternary *Helicosphaera sellii* to *Calcidiscus macintyrei* Zone. A. Abundance of diatom valves per gram sediment (logarithmic scale). B. Percent of CH₃COOH-insoluble residue >2 µm of the bulk sediment. C. Downcore occurrence of diatoms plotted vs. age. The age assignment is based on the calcareous nannofossil stratigraphy (Premoli Silva, Haggerty, Rack, et al., 1993; see also various papers in this volume). The hatched part of the columns represents a hiatus. D. Downcore occurrence of diatoms plotted vs. depth.

eros messanensis, among others, can be common to frequent. The diversity at Site 880 is relatively high, reaching as many as 33 species per 100 counted specimens (Table 6). Overall, the assemblages have a temperate character. Sporadically, a fragment or valve of a preferably tropical-subtropical species such as *Asteromphalus elegans* or *Asterolampra marylandica* occurs. But cosmopolitan and temperate species are dominant. The presence of temperate endemic northern Pacific species such as *Neodenticula kamtschatica* and *N. seminae* is quite conspicuous, as are the presence of species characteristic for higher productivity shelf and shelf-edge environments, such as *Cyclotella litoralis*, species of the genus *Chaetoceros*, *Nitzschia bicapitata*, *Odontella aurita*, *Stephanopyxis turris*, and *Paralia sulcata*.

Concerning the biostratigraphy based on diatoms, the *Neodenticula seminae* Zone and the *Rhizosolenia curvirostris* Zone can be identified within the upper 7 m of the cores (Table 6). The boundary between the two zones lies between Samples 144-880A-1H-1, 62 cm, and -1H-1, 64 cm. It was determined using the last occurrence of *Thalassiosira nidulus* (0.28–0.39 Ma; Barron, 1992). The marker species *Simonsenella curvirostris* is absent from the samples studied. The base of the *R. curvirostris* Zone cannot be determined because the diatom preservation below 7 m is too poor.

Some reworking of older material was also found at Site 880. Single valves of early Oligocene (*Pyxilla reticulata*), Miocene (*Denticulopsis lauta*), and early Pliocene (*Rossiella tatsunokuchiensis*) diatom species were encountered in different samples (Table 6).

Latitudinal Trends

Tracing latitudinal trends is complicated by the fact that, because of erosion or sediment loss during coring, the uppermost Quaternary stratigraphic intervals are not represented at some sites. For example, the *Emiliania huxleyi* Acme Zone, within which the best preserved siliceous microfossil assemblages can be expected, was recovered only at Sites 871 and 880; at Sites 873 and 878 no upper Quaternary sediments were recovered. Nevertheless, for two stratigraphic intervals of the Quaternary, a south-north comparison was possible: (1) for the *Emiliania huxleyi* Zone, which was cored at Sites 871, 872, and 880; and (2) for the *Helicosphaera sellii* to *Calcidiscus macintyrei* zones, an interval recovered at Sites 871, 872, 873, 878, and 880.

Diatom abundance is low at the three southern subtropical sites, and their preservation is poor. Their abundance (Fig. 8) increases by approximately 1 order of magnitude at about 27°N, on MIT Guyot (Site 878), and by up to 3 orders of magnitude at the northernmost site, on Takuyo-Daisan Guyot (Site 880). Here diatoms are better preserved and a higher diversity of their assemblages is found (Table 6). Unfortunately, at this site, the base of diatom occurrence cannot be determined because Hole 880 was drilled only 18 m into the upper Pliocene. Diatoms are still present at that depth, although at a lower diversity, and it is the more dissolution-resistant species that occur at these core depths.

Table 11. Logarithmic abundance of major siliceous microfossil groups, pollen, and vesicular ash, and percent abundance of CH_3COOH -insoluble residue >2 μm in samples from Hole 880A.

Core, section, interval (cm)	Core depth (cm bsf)	CH_3COOH -insoluble residue (%)	Siliceous sponge spicules						Freshwater diatoms			Pollen
			Vesicular ash	<i>Actiniscus</i>	Silicoflagellates	Marine diatoms	Radiolarians	Archaeomonadaceae	Freshwater diatoms	Phytoliths		
144-880A-												
1H (top)	0-1	ND	6.6	—	5.1	6.4	5.5	5.0	4.5	5.0	5.1	—
1H-1, 2	2	ND	6.2	4.2	4.6	5.7	4.5	5.0	—	4.3	4.1	3.8
1H-1, 4	4	ND	6.1	—	—	4.6	—	3.9	—	3.6	—	3.6
1H-1, 6	6	ND	6.0	—	1.4	3.4	1.4	3.0	3.9	3.0	3.0	2.4
1H-1, 8	8	ND	6.2	—	—	4.5	3.7	3.7	4.0	4.0	3.7	—
1H-1, 10	10	ND	6.0	—	—	4.2	3.7	3.9	—	3.4	—	3.4
1H-1, 12	12	ND	5.7	—	3.3	4.4	—	3.8	3.6	4.0	3.6	3.3
1H-1, 14	14	ND	5.9	—	—	4.2	3.3	3.3	3.3	4.1	3.3	3.8
1H-1, 18	18	ND	5.9	—	3.1	4.1	—	3.8	4.1	3.8	3.8	3.1
1H-1, 20	20	ND	6.1	—	ND	3.5	—	3.8	4.5	3.5	3.5	—
1H-1, 22	22	ND	6.2	ND	—	4.8	4.1	4.1	4.2	3.4	4.0	—
1H-1, 24	24	ND	6.2	—	3.2	4.9	3.9	4.1	3.5	3.5	3.8	3.2
1H-1, 26	26	ND	6.1	—	—	4.9	4.3	4.1	4.1	4.3	—	—
1H-1, 28	28	ND	6.2	—	—	5.0	4.3	4.1	4.0	4.1	4.0	3.7
1H-1, 30	30	ND	6.0	—	3.8	5.0	4.4	—	4.2	3.8	4.1	—
1H-1, 32	32	ND	6.3	—	4.3	5.2	4.8	4.4	3.9	3.9	4.6	—
1H-1, 34	34	ND	5.8	—	3.1	4.2	3.4	4.0	4.4	3.1	3.6	—
1H-1, 36	36	ND	6.1	—	3.6	4.5	4.1	4.1	4.2	—	3.9	—
1H-1, 38	38	ND	5.8	—	—	3.9	2.2	3.5	4.6	—	—	3.5
1H-1, 40	40	ND	6.0	—	—	4.5	—	4.0	4.3	4.0	—	4.0
1H-1, 42	42	ND	6.0	—	—	5.1	4.5	4.2	—	4.2	—	—
1H-1, 44	44	ND	6.2	—	3.9	5.4	4.6	4.2	4.2	4.6	—	3.9
1H-1, 46	46	ND	6.4	—	—	5.3	4.3	4.6	4.3	4.3	4.0	—
1H-1, 48	48	ND	6.4	—	3.3	4.7	4.0	4.2	4.7	3.8	3.3	3.3
1H-1, 50	50	ND	6.3	—	3.8	5.1	4.3	4.1	4.1	4.6	—	—
1H-1, 53	53	ND	6.9	—	—	3.0	—	4.0	—	—	—	—
1H-1, 56	56	ND	6.3	—	3.7	5.2	4.4	4.2	—	4.2	4.0	—
1H-1, 58	58	ND	5.8	3.9	3.9	5.2	4.6	4.3	4.3	4.1	3.9	3.9
1H-1, 60	60	ND	6.4	3.9	—	5.0	4.4	4.2	—	4.5	4.5	3.9
1H-1, 62	62	ND	6.1	3.7	—	4.9	4.0	4.4	4.0	3.7	4.4	—
1H-1, 64	64	ND	6.2	—	—	4.7	4.1	4.0	4.1	3.4	4.0	—
1H-1, 66	66	ND	6.1	2.9	2.9	4.9	4.4	4.4	4.9	3.2	3.8	2.9
1H-1, 73-74	73-74	61.8	6.4	—	5.1	6.2	5.1	4.8	4.6	4.4	4.4	4.1
1H-1, 99-100	99-100	81.1	6.3	3.7	5.0	5.7	4.8	4.6	3.7	—	3.7	—
1H-1, 120	120	ND	6.3	—	—	4.0	—	—	4.0	—	—	4.0
1H-1, 129-130	129-130	80.6	6.5	—	4.8	5.7	4.4	4.1	4.3	3.7	3.7	—
1H-1 (base)	149-150	ND	6.0	—	—	4.8	—	3.7	4.1	3.9	3.7	3.4
1H-2 (???)	159-160	59	6.3	—	4.3	5.5	4.3	4.4	4.4	3.5	4.0	—
1H-2, 59-60	209-210	83.8	6.7	3.4	4.4	5.4	4.5	4.2	4.5	3.9	—	—
1H-2, 128-129	278-279	77.8	6.3	—	3.3	4.6	4.4	3.9	3.3	—	3.6	3.6
1H-2 (base)	299-300	ND	6.3	—	—	4.2	—	3.7	4.0	4.2	3.3	—
1H-3, 49-50	349-350	83.1	6.5	3.7	4.7	5.7	4.3	3.7	4.4	4.0	—	3.7
1H-3, 109-110	409-410	72.8	6.1	—	3.6	4.4	3.9	3.7	3.6	—	—	—
1H-3 (base)	449-450	ND	6.2	—	—	4.0	3.7	3.7	—	—	—	3.7
1H-4, 50-51	500-501	60.5	6.2	—	—	3.9	—	3.6	3.2	—	—	3.2
1H-4, 108-109	558-559	69.9	6.7	—	4.0	5.3	4.7	4.0	4.3	4.0	4.0	—
1H-4, 140-150	590-600	68.3	6.8	—	3.3	4.4	4.4	3.8	—	3.8	3.6	3.6
1H-5, 48-49	648-649	80.2	6.3	—	4.0	5.6	4.8	3.8	4.1	—	—	—
1H-5, 109-110	709-710	49.6	5.9	—	—	4.6	3.4	3.1	3.4	3.6	3.4	—
1H-6, 53-54	803-804	72.5	6.2	—	—	3.4	—	—	4.2	3.4	3.4	3.4
1H-6, 108-109	858-859	66.8	6.2	—	—	4.1	—	3.7	4.1	—	—	3.4
2H-CC (upper)	864-865	ND	6.6	—	—	4.3	3.3	3.3	3.3	3.8	3.3	—
2H-CC (lower)	885-886	ND	6.2	—	—	4.3	—	3.7	4.0	3.5	3.2	—
2H-1, 48-49	938-939	52.8	6.3	—	—	4.1	—	—	4.3	—	—	—
2H-1, 108-109	998-999	46.2	6.7	—	3.3	4.3	3.8	3.7	—	3.4	—	3.4
2H-2, 49-50	1074-1075	46.7	6.4	—	3.5	4.4	—	4.2	4.2	3.8	3.8	3.8
2H-2, 109-110	1134-1135	79.3	6.4	—	—	4.2	—	3.5	4.4	3.8	—	—
2H-3, 48-49	1223-1224	79.9	6.1	—	—	4.1	3.2	3.2	3.8	—	—	—
2H-3, 109-110	1284-1285	78.3	6.1	—	—	4.1	—	—	4.6	3.5	—	—
2H-4, 49-50	1374-1375	64.8	6.6	—	—	3.9	—	3.9	4.5	3.6	3.6	—
2H-4, 107-108	1432-1433	83.5	6.5	—	—	3.9	—	—	3.9	—	—	—
2H-5, 49-50	1524-1525	79.2	6.2	—	—	3.1	—	3.4	4.5	3.4	ND	3.4
2H-5, 108-109	1583-1584	75.3	6.3	—	—	3.7	—	3.7	4.2	3.9	—	—
2H-5, 140-150	1615-1625	68.9	6.6	—	—	3.9	3.2	3.2	3.7	3.7	3.5	3.2
2H-6, 49-50	1674-1675	62.8	6.3	—	—	3.7	—	3.9	3.9	—	3.4	3.4
2H-6, 109-110	1734-1735	66.9	6.4	—	—	3.6	3.3	3.7	3.6	3.6	—	3.3
2H-6, 142-150	1767-1775	64.2	6.9	—	—	3.4	2.8	3.3	3.3	3.8	—	—
2H-7, 49-50	1824-1825	79.9	6.6	—	—	3.9	3.1	4.0	4.0	3.7	—	—
2H-CC	1830-1831	ND	6.2	—	—	3.9	—	3.4	4.0	3.5	3.4	3.1

Note: ND = not determined.

A correlation of diatom abundance (Fig. 8A) and preservation with the amount of CH_3COOH -insoluble residue >2 μm (Fig. 8B) and an inverse correlation with the calcium carbonate content of the sediment can be recognized.

The CH_3COOH -insoluble residue predominantly consists of wind-transported volcanic ash. Its amount in the highly calcareous, tropical-subtropical, Quaternary pelagic sediments is minimal. At approximately 27°N, ash content fluctuates during the early Quaternary

around 20%, and around 70% (between 40% and 90%) at approximately 34°N. Such quantities, especially if it is vesicular ash as at Site 880, can influence the preservation of biogenic silica, although the higher clay content possibly plays a role as well.

Another influential factor is the increased primary productivity at these two northern locations, especially off Japan (Koblenz-Mishke et al., 1970). That productivity at Site 880 is higher than at the other sites drilled further south is shown (1) by the fact that, in addition to

an increase in Pleistocene sedimentation rates (at Sites 871–873, they lie between 1 and 10 m/m.y. and at Site 880 between 10 and 20 m/m.y.; cf. Erba, this volume) within these sediments, the abundance of planktonic marine diatoms increases up to 3 orders of magnitude from the southern sites to Site 880; and (2) by the presence of *Spermatogonia* sp. and *Actiniscus* sp. (Table 6), which also indicates more fertile surface waters. The increased primary productivity in this region was also illustrated by the frequency of fishing boats met during the cruise: none were seen until we approached Takuyo-Daisan. In addition, the benthic life on this guyot is relatively rich. On the more southern guyots, no macroorganisms were seen with the camera, only barren rock or pelagic sediments in part with ripples on the surface. On Takuyo-Daisan, the camera showed corals, sponges, and possibly crinoids. Such benthic life is proof that sufficient nutrients are reaching the seafloor.

If one plots the age of the sediment obtained from calcareous nanofossil stratigraphy (Premoli Silva, Haggerty, Rack, et al., 1993) down to which diatoms were found at each site, the result is that the lowest occurrence of diatoms in the sediments is increasingly older toward the north (Fig. 8C). However, because a stratigraphic interval is missing at the top of the subtropical sites, the core depth down to which diatoms are preserved within these sites does not show such a clear trend (Fig. 8D).

Pore-water analysis suggests that bottom-water concentrations perpetuate throughout the sandy foraminifer-nanofossil oozes of the pelagic cap sediments and no special conditions in the near-surface sediments can be recognized.

SUMMARY

The siliceous microfossil assemblages at all of the Leg 144 sites studied are influenced by silica dissolution and diagenesis. This influence is especially severe in the highly calcareous sediments of the tropical-subtropical pelagic sediments. It is weakest in the less calcareous pelagic sediments off Japan, where vesicular ash constitutes a major component of the sediment and may help to protect siliceous microfossils from dissolution, and where the accumulation rates of biosiliceous components are higher.

The types of siliceous sponge spicules from desmospongia and hexactinellid sponges living on the guyots, in water depths of 1000–1500 m, are documented.

Only at Site 880, on Takuyo-Daisan Guyot off Japan, are the preservation and diversity within the diatom assemblages sufficient to allow determination of the diatom biostratigraphic zones established for the North Pacific. The last occurrence datum of *Thalassiosira nidulus* indicates that the upper 62 cm of the Neogene pelagic sediments would belong to the *Neodenticula seminae* Zone and the sediments below, down to 7 m, to the *Rhizosolenia curvirostris* Zone.

In spite of the poor preservation of diatoms in most samples analyzed, geographical restrictions of some species were evident. For example, *Asteromphalus sacrophagus* var. *obovatus* was only found at the southernmost location (Site 871), whereas *Neodenticula seminae* and *N. kamtschatica* occur persistently only at the northernmost site (880) and with single sporadic occurrences at Hole 872A.

The increase of diatom abundance from the three sites at 5–15°N, where it is low, to higher values at 27°N and even higher at 34°N, off Japan, is largely a result of higher primary productivity in the belt influenced by west winds.

The abundance of freshwater diatoms, phytoliths, and pollen as well as of vesicular ash on Takuyo-Daisan Guyot document a high eolian component in the pelagic sediments east of Japan. Close sample spacing from the upper 70 cm recovered at Hole 880A allows the recognition of climate-related changes in abundance of siliceous microfossils. Because these abundance changes show the same trend as the abundance of volcanic ash, dilution by windblown terrigenous

components and possibly preservational effects are thought to be responsible for the abundance changes.

TAXONOMIC REFERENCES FOR MARINE DIATOM SPECIES

In the following, all diatom species are listed with reference either to the original or to a more modern description that was used as a basis for species determination in this study. A few remarks are made only where species determination is not unanimously handled in the recent literature, and only those species are illustrated on Plate 1.

Genus *Actinocyclus* Ehrenberg (1837)
Actinocyclus curvatus Janisch in Schmidt (1878)
 Pl. 1, Fig. 7

Description. A. Schmidt (1874–1900), pl. 57, fig. 31

Remarks. The valves are convex. The areolae decrease in size toward the margin and form segments, the longest ray of each being clearly curved.

Actinocyclus ellipticus Grunow in Van Heurck (1883)

Description. Van Heurck (1883), pl. 124, fig. 10; Hustedt (1930), p. 533, fig. 303

Actinocyclus ingens Rattray forma *planus* Whiting et Schrader (1985)

Description. Whiting and Schrader (1985), p. 74, pl. 3, fig. 12

Actinocyclus octonarius Ehrenberg (1838)

Description. Ehrenberg (1838), p. 173, pl. 21, fig. 7; Hendey (1964), p. 83, pl. 24, fig. 3

Actinocyclus tenellus (Brébisson) Andrews (1976)

Description. Andrews (1976), p. 14, pl. 3, figs. 8, 9

Actinocyclus sp. 1
 Pl. 1, Fig. 6

Description. The valve outline is round. The valve face is flat and densely covered with areolae arranged in fascicles and slightly decreasing in size from the center (10–14 areolae in 10 µm) toward the margin (12–16 areolae in 10 µm). There are 7–8 marginal labiate processes. They are located at the end of the longest row of each fascicle. The areolae on the mantle are small (approximately 15 in 10 µm) and arranged in a dense quincunx pattern, distinctly offset from the pattern on the valve face.

Genus *Actinoptychus* Ehrenberg (1841)
Actinoptychus senarius (Ehrenberg) Ehrenberg (1838)

Description. Hendey (1964), p. 95, pl. 23, figs. 1–2

Actinoptychus splendens (Shadboldt) Ralfs in Pritchard (1861)

Description. Hustedt (1930), p. 478, fig. 265

Genus *Asterolampra* Ehrenberg (1845)
Asterolampra marylandica Ehrenberg (1845)

Description. Ehrenberg (1845), p. 76, fig. 10

Genus *Asteromphalus* Ehrenberg (1845)
Asteromphalus elegans Greville (1859)

Description. Greville (1859), p. 161, pl. 7, fig. 6

Asteromphalus hiltonianus (Greville) Ralfs in Pritchard (1861)

Description. Pritchard (1861), p. 837; Greville (1860), p. 117, pl. 4, fig. 15

Asteromphalus hookeri Ehrenberg (1844c)

Description. Ehrenberg (1845), pl. 35A/21, fig. 2

Asteromphalus ornithopus Karsten (1907)
 Pl. 4, Fig. 6

Description. Karsten (1907), p. 90, pl. 8, fig. 13

Asteromphalus robustus Castracane (1875)

Description. Castracane (1875), p. 393, pl. 6, fig. 5

Asteromphalus sarcophagus Wallich forma *obovatus* Thorington-Smith (1970)

Description. Thorington-Smith (1970), pp. 821–822, pl. 1, fig. 3

Asteromphalus sarcophagus Wallich forma *sarcophagus* Thorington-Smith (1970)

Description. Thorington-Smith (1970), pp. 820–821, pl. 1, fig. 1

Genus *Aulacoseira* Thwaites (1848)

Pl. 2, Fig. 20

Remark. The species within this genus all are freshwater species and were not determined to the species level.

Genus *Azpeitia* M. Peragallo in Tempère et Peragallo (1912)

Azpeitia africana (Janisch ex Schmidt) Fryxell et Watkins (1986)

Description. Fryxell, Sims, and Watkins (1986), pp. 22–23, figs. 22–23,

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Azpeitia neocrenulata (Van Landingham) Fryxell et Watkins (1986)

Description. Fryxell, Sims, and Watkins (1986), pp. 18–19, figs. 16, 30

Azpeitia nodulifer (Schmidt) Fryxell et Sims (1986)

Description. Fryxell, Sims, and Watkins (1986), pp. 19–20, figs. 17, 18,

30

Azpeitia tabularis (Grunow) Fryxell et Sims (1986)

Description. Grunow (1884), p. 86; Fryxell, Sims, and Watkins (1986), pp. 16–18, figs. 14–15, 30

Genus *Bacteriastrum* Shadbolt (1854)

Bacteriastrum sp.

Remark. The only valve found (at Site 880) is incomplete. No species determination was possible.

Genus *Chaetoceros* Ehrenberg (1844a)

Chaetoceros messanensis Castracane (1875)

Pl. 1, Fig. 1

Description. Castracane (1875), p. 394, pl. 1, fig. 1a; Hustedt (1920), pl. 322, figs. 4, 7c; pl. 325, fig. 3

Chaetoceros sp.

Remark. Relatively robust valve fragments of other *Chaetoceros* species with the setae generally broken above their basal part were found sporadically but were not determined to the species level.

Genus *Coccoeis* Ehrenberg (1838)

Coccoeis placentula Ehrenberg (1838)

Pl. 2, Fig. 19

Description. Ehrenberg (1841), pl. 1/1, fig. 10, pl. 1/2, fig. 24; Hustedt (1933), p. 347, fig. 802a–b

Genus *Coscinodiscus* Ehrenberg (1838)

Coscinodiscus argus Ehrenberg (1839)

Description. Hasle and Sims (1986b), pp. 308–310, 312, 316, figs. 1–7, 33–34

Coscinodiscus gigas Ehrenberg (1841)

Description. Ehrenberg (1854), pl. 18, fig. 34.

Coscinodiscus marginatus Ehrenberg (1841)

Description. Ehrenberg (1841), p. 142; Ehrenberg (1854), pl. 18, fig. 44, pl. 33, figs. 12–13, pl. 38B/22, fig. 8

Coscinodiscus pacificus (Grunow) Rattray (1890)

Pl. 4, Fig. 1

Description. Rattray (1890), p. 563

Coscinodiscus radiatus Ehrenberg (1840)

Pl. 1, Fig. 8

Description. Hasle and Sims (1986a), pp. 310–314, 317, figs. 8–32, 35–

39

Genus *Cyclotella* Kützing (1833)

Cyclotella litoralis Lange et Syvertsen (1989)

Pl. 1, Figs. 2–3

Description. Lange and Syvertsen (1989), pp. 343–346, figs. 1–30

Cyclotella radiosua (Grunow) Lemmermann (1900)
Pl. 2, Fig. 22

Description. Round and Håkansson (1992), pp. 114–116, figs. 9–12, 16, 32–35

Remark. *C. radiosua* is a freshwater species.

Genus *Denticulopsis* Simonsen (1979)
Denticulopsis lauta (Bailey) Simonsen (1979)
Pl. 1, Fig. 12

Description. Yanagisawa and Akiba (1990), pp. 235–236, pl. 2, figs. 6–8, 15; pl. 5, figs. 1–3; pl. 9, fig. 1

Genus *Diploneis* Ehrenberg (1840)
Diploneis bombus Ehrenberg (1844b)
Pl. 4, Fig. 4

Description. Hustedt (1959), pp. 704–709, fig. 1086.

Genus *Ethmodiscus* Castracane (1866)

Remarks. As only relatively small fragments of *Ethmodiscus* valves were found no attempt was made to determine the species.

Genus *Fragilariopsis* Hustedt (1913) in A. Schmidt (1874–1900)
Fragilariopsis doliolus (Wallich) Medlin et Sims (1993)

Description. Medlin and Sims (1993), p. 332, figs. 5–12

Genus *Hemidiscus* Wallich (1860)
Hemidiscus cuneiformis Wallich (1860)

Description. Wallich (1860), p. 42, pl. 2, figs. 3–4; Hustedt (1930), p. 904, fig. 542e; Fryxell, Sims, and Watkins (1986), p. 25, fig. 26

Hemidiscus cuneiformis Wallich var.
Pl. 1, Fig. 17

Remark. This variety differs from the species by its nearly elliptical shape and the much finer areolation.

Hemidiscus sp. 1
Pl. 1, Fig. 5

Description. This species is similar in shape to *Hemidiscus karstenii* Jousé in Jousé et al. (1962), but differs from it because of its finer and denser areolation. Also, the size of the areolae decreases only slightly from the valve center toward the margin.

Genus *Navicula* Bory (1822)
Navicula directa (Wm. Smith) Ralfs in Pritchard (1861)

Description. Hendey (1964), p. 202; A. Schmidt (1874–1900), pl. 47, fig. 5

Genus *Neodenticula* Akiba et Yanagisawa (1986)
Neodenticula kamtschatica (Zabelina) Akiba et Yanagisawa (1986)
Pl. 1, Figs. 19–21

Description. Yanagisawa and Akiba (1990), pp. 259–261, pl. 7, figs. 27–37
Neodenticula seminae (Simonsen et Kanaya) Akiba et Yanagisawa (1986)
Pl. 1, Figs. 13–15

Description. Yanagisawa and Akiba (1990), pp. 263–264, pl. 7, figs. 45–

49

Genus *Nitzschia* Hassall (1845)
Nitzschia bicapitata Cleve (1900)

Description. Cleve 1900, p. 933, fig. 12

Nitzschia braarudii Hasle (1960)

Description. Hasle (1960), p. 22, pl. 7, figs. 58–63

Nitzschia interruptestriata Simonsen (1974)
Pl. 1, Fig. 11

Description. Simonsen (1974), p. 52, pl. 36, figs. 9–11; pl. 37, pl. 38, figs. 1–7

Nitzschia kolaczekii Grunow (1867)

Description. Grunow (1877), p. 173, pl. 194, fig. 10

Nitzschia marina Grunow in Cleve and Grunow (1880)

Description. Grunow in van Heurck (1881), pl. 57, figs. 26–27

- Nitzschia sicula* (Castracane) Hustedt (1958)
- Description.** Hustedt (1958), p. 180
- Genus *Odontella* Agardh (1832)
 - Odontella aurita* Agardh (1832)
- Description.** Hustedt (1930), pp. 846–849, fig. 501
- Genus *Paralia* Heiberg (1863)
 - Paralia sulcata* (Ehrenberg) Cleve (1873)
- Description.** Hustedt (1930), pp. 276–279, figs. 118–120
- Genus *Planktoniella* Schütt (1893)
 - Planktoniella sol* (Wallich) Schütt (1893)
 - Pl. 4, Fig. 3
- Description.** Gerloff (1970), pp. 203–215
- Genus *Pseudotriceratum* Grunow (1884)
 - Pseudotriceratum cinnamomeum* (Greville) Grunow (1884)
 - Pl. 1, Fig. 4
- Description.** Grunow (1884), p. 83; Greville (1863), p. 232, pl. 9, fig. 12
- Genus *Pyxilla* Greville (1865)
 - Pyxilla reticulata* Grove et Sturt (1887)
- Description.** Fenner (1985), pp. 735–736, pl. 11, figs. 6–10
- Remark.** The fragment found at Site 880 must be reworked from Paleogene sediments, probably lower Oligocene.
- Resting Spore
 - Pl. 4, Fig. 7
- Remark.** As resting spores of diatoms were found only relatively rarely, only two groups were differentiated, the *Xanthiopxis* type resting spores and other resting spores. But they were not determined to the species level.
- Genus *Raphoneis* Ehrenberg (1844c)
 - Raphoneis surireloides* Simonsen (1974)
 - Pl. 4, Fig. 2
- Description.** Simonsen (1974), pp. 35–36, pl. 23, figs. 2–8
- Genus *Rhizosolenia* Ehrenberg (1841)
 - Rhizosolenia bergenii* H. Peragallo (1892)
- Description.** Hustedt (1930), pp. 575–577, fig. 327
- Rhizosolenia hebetata* Bailey (1856)
- Description.** Grunow (1884), p. 44, pl. 5, figs. 48–50
- Rhizosolenia setigera* Brightwell (1858)
- Description.** van Heurck (1880–1885), p. 194, pl. 78, figs. 6–8
- Rhizosolenia styliformis* Brightwell (1858)
- Description.** van Heurck (1880–1885), p. 194, pl. 78, figs. 1–5; pl. 79, figs. 1–2
- Rhizosolenia styliformis* var. *longispina* Hustedt (1914)
- Description.** Hustedt (1914), pl. 316, figs. 5–7, 12; Hustedt (1930), p. 586, fig. 334
- Genus *Roperia* Grunow in van Heurck (1883)
 - Roperia tessellata* (Roper) Grunow in van Heurck (1883)
- Description.** Fryxell, Sims, and Watkins (1986), p. 24, fig. 25
- Genus *Rossiella* Desikachary et Maheshwari (1958)
 - Rossiella tatsunokuchiensis* (Koizumi) Gersonde et Schrader (1984)
- Description.** Koizumi (1972), p. 349, pl. 42, figs. 3–4
- Genus *Rouxia* Brun et Héribraud in Héribraud (1893)
 - Rouxia californica* M. Peragallo in Tempère and Peragallo (1910)
- Description.** Tempère and Peragallo (1910), p. 245, nos. 468–469
- Genus *Simonseniella* Fenner (1991)
 - Simonseniella barboi* (Brun) Fenner (1991)
- Synonym.** *Proboscia barboi* (Brun) Jordan and Priddle (1991), p. 56, figs. 1–2
- Description.** Fenner (1991), p. 108, pl. 3, figs. 1, 3
- Genus *Stellarima* Hasle et Sims (1986a)
- Stellarima microtrias* (Ehrenberg) Hasle et Sims (1986b)
- Description.** Hasle and Sims (1986a), p. 111, figs. 22–27
- Genus *Stephanopyxis* Ehrenberg (1844d)
 - Stephanopyxis turris* (Greville et Arnott) Ralfs in Pritchard (1861)
- Description.** Grunow in van Heurck (1882), pl. 83 ter., fig. 12
- Genus *Thalassionema* Grunow in van Heurck (1881)
 - Thalassionema bacillaris* (Heiden) Kolbe (1955)
- Description.** Hasle and Mendiola (1967), pp. 107–125, figs. 1–4, 6–8, 18–19, 22–26; Hallegraaff (1986), pp. 60–62, fig. 5; Simonsen (1992), p. 25, pl. 22, figs. 1–6
- Thalassionema nitzschiooides* Grunow in van Heurck (1881)
- Description.** Hallegraaff (1986), pp. 58–60, figs. 1–4
- Thalassionema nitzschiooides* var. *obtusa* (Grunow) M. Peragallo (1903)
- Description.** Hasle (1960), p. 18, figs. 5b, 5e
- Thalassionema nitzschiooides* var. *parva* Heiden et Kolbe (1928)
- Description.** Hasle (1960), p. 18, fig. 5a
- Genus *Thalassiosira* Cleve (1873)
 - Thalassiosira anguste-lineata* (A. Schmidt) Fryxell et Hasle (1977)
- Description.** Fryxell and Hasle (1977), pp. 73–75, figs. 22–34
- Thalassiosira antiqua* (Grunow) Cleve-Euler (1941)
- Description.** Cleve-Euler (1951), p. 72, fig. 119
- Thalassiosira cf. decipiens* (Grunow) Jørgensen (1905)
 - Pl. 1, Fig. 18
- Description.** Hustedt (1930), pp. 322–323, fig. 158
- Thalassiosira eccentrica* (Ehrenberg) Cleve (1903)
- Description.** Fryxell and Hasle (1972), pp. 300–312, figs. 1–18
- Thalassiosira ferelineata* Hasle et Fryxell (1977)
- Description.** Hasle and Fryxell (1977), pp. 26–28, figs. 46–53
- Thalassiosira gravida* Cleve (1896)
- Description.** Cleve (1896), p. 12, pl. 2, figs. 14–16; Hustedt (1930), p. 325, fig. 161; Syvertsen (1979), pp. 52–54
- Thalassiosira latimarginata* Makarova (1975)
- Synonym.** *Thalassiosira trifulta* Fryxell in Fryxell and Hasle (1979) pp. 16–19, figs. 1–24
- Description.** Makarova (1975), p. 150, figs. 3–4
- Thalassiosira leptopus* (Grunow) Hasle et Fryxell (1977)
- Description.** Hasle and Fryxell (1977), pp. 20–22, figs. 1–14, 94–96
- Thalassiosira lineata* Jousé (1968)
- Description.** Hasle and Fryxell (1977), pp. 22–23, figs. 15–25
- Thalassiosira nidulus* (Tempère et Brun) Jousé (1961)
 - Pl. 4, Fig. 5
- Description.** Jousé (1961), p. 63, pl. 3, figs. 4–5
- Thalassiosira oestrupii* (Ostenfeld) Proschkina-Lavrenko var. *oestrupii* Jousé (1968)
- Description.** Jousé (1968), pp. 13–15, pl. 1, figs. 3–7
- Thalassiosira oestrupii* var. *plana* Jousé (1968)
- Description.** Jousé (1968), p. 15, pl. 1, figs. 8, 9
- Thalassiosira symmetrica* Fryxell et Hasle (1972)
- Description.** Fryxell and Hasle (1972), p. 312, figs. 37–46
- Thalassiosira* sp. 1
 - Pl. 1, Fig. 16
- Remarks.** Only single fragments of this species were found at Sites 873 and 880. The valves are convex. The areolae decrease in size toward the margin and form segments. The margin is coarsely striate (10–11 striae in 10 µm).

Genus *Thalassiothrix* Cleve et Grunow (1880)
Thalassiothrix longissima Cleve et Grunow (1880) in Cleve et Möller (1878)

Description. Hasle and Semina (1987), pp. 177–181, figs. 1–25

Genus *Xanthiopyxis* Ehrenberg (1944d)
 Pl. 4, Fig. 8

Remarks. No species were differentiated.

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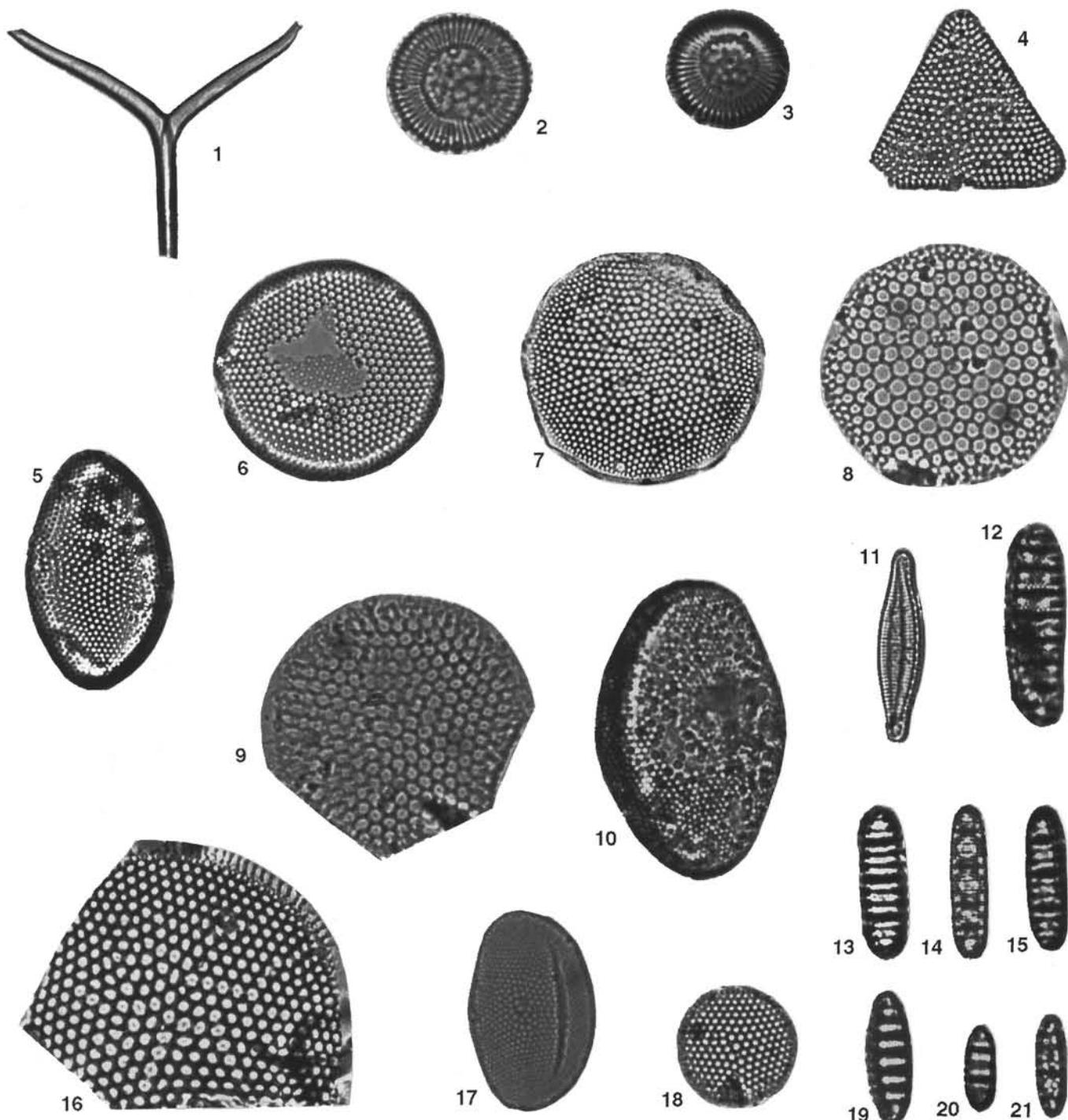


Plate 1. Marine planktonic diatoms in Quaternary sediments. Magnification 1500 \times , except for Figure 7, which is at 975 \times . 1. *Chaetoceros messanensis*, Sample 144-880A-1H-1, 73–74 cm. 2–3. *Cyclotella littoralis*; (2) Sample 144-880A-1H-1, 60 cm; (3) Sample 144-880A-1H-1, 24 cm. 4. *Pseudotriceratium cinnamomeum*, Sample 144-880A-1H-1, 46 cm. 5. *Hemidiscus* sp. 1, Sample 144-880A-1H-2, 59–60 cm. 6. *Actinocyclus* sp. 1, Sample 144-880A-1H-2, 9–10 cm. 7. *Actinocyclus curvatus*, Sample 144-873B-1H-1, 30–34 cm. 8. *Coscinodiscus radiatus*, Sample 144-880A-1H-1, 60 cm. 9. *Thalassiosira* sp., Sample 144-880A-1H-1, 42 cm. 10. *Hemidiscus?* sp., Sample 144-880A-1H-1, 30 cm. 11. *Nitzschia interruptestriata*, Sample 144-880A-1H-1, 129–130 cm. 12. *Denticulopsis lauta*, Sample 144-880A-1H-1, 48 cm. 13–15. *Neodenticula seminae*, 13; Sample 144-880A-1H-1, base. 14. Sample 144-880A-1H-1, 18 cm. 15. Sample 144-880A-1H-1, 24 cm. 16. *Thalassiosira* sp. 1, Sample 144-880A-1H-1, 24 cm. 17. *Hemidiscus cuneiformis* var., Sample 144-880A-1H-1, 32 cm. 18. *Thalassiosira* aff. *decipiens*, Sample 144-880A-1H-1, 73–74 cm. 19–21. *Neodenticula kamtschatica*; (19) Sample 144-880A-1H-1, 14 cm; (20) Sample 144-880A-1H-1, 36 cm; (21) Sample 144-880A-1H-1, 18 cm.

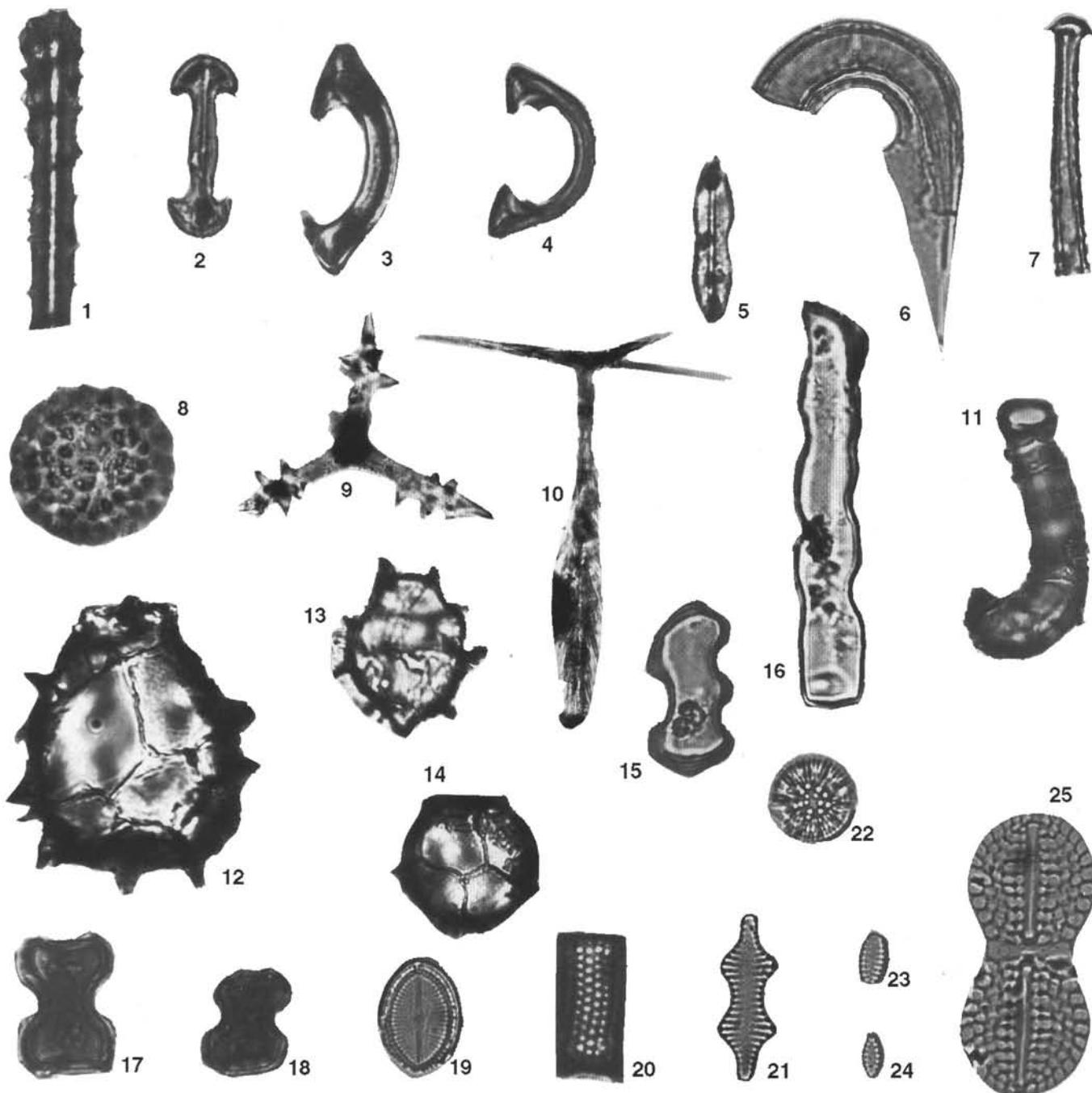


Plate 2. Magnification 1500 \times , except for Figures 1, 9, 11, and 13 at 975 \times ; Figures 8 and 14 at 625 \times ; and Figure 10 at 300 \times . Figures 1-10 represent siliceous sponge spicules and microscleres. 1. *Acanthostyle* (m), Sample 144-880A-1H-1, 48 cm. 2-4. *Arcuate isoschele* (m); (2) front view, Sample 144-880A-1H-1, 36 cm; (3) side view, Sample 144-880A-1H-1, 62 cm; (4) side view, Sample 144-880A-1H-1, 66 cm. 5. Unknown type (m), Sample 144-880A-1H-2, 59-60 cm. 6. *Diancistra?* (m), Sample 144-880A-1H-1, 62 cm. 7. Fragment of forceps (m), Sample 144-880A-1H-2, 9-10 cm. 8. *Sterosphaeraster* (m), Sample 144-873B-1H-1, 30-34 cm. 9. *Acanthotetraxon* (m), Sample 144-873B-1H-1, 30-34 cm. 10. *Pentactinipinul* (m), Sample 144-878A-1H-1 (top). 11. Organic walled tube, Sample 144-880A-1H-1, 66 cm. Figures 12-14 are dinoflagellates. 12. Sample 144-872A-1H-5, 39-43 cm. 13. Sample 144-872A-1H-4, 140-144 cm. 14. Sample 144-872A-1H-CC. Figures 15-18 are phytoliths. 15. Sample 144-880A-1H-1, 62 cm. 16. Sample 144-880A-2H-2, 49-50 cm. 17. Sample 144-880A-1H-1, 6 cm. 18. Sample 144-880A-1H-1, 46 cm. Figures 19-25 are freshwater diatoms. 19. *Cocconeis placentula*, Sample 144-880A-1H-1, 36 cm. 20. *Aulacoseira* sp., Sample 144-880A-1H-1, 26 cm. 21. *Fragilaria* sp., Sample 144-880A-1H-1, 18 cm. 22. *Cyclotella radiosa*, Sample 144-880A-1H-1, base. 23, 24. *Fragilaria* sp.; (23) Sample 144-880A-1H-1, 58 cm; (24) Sample 144-880A-1H-1, 60 cm. 25. *Diploneis* sp., Sample 144-880A-1H-1, 66 cm.

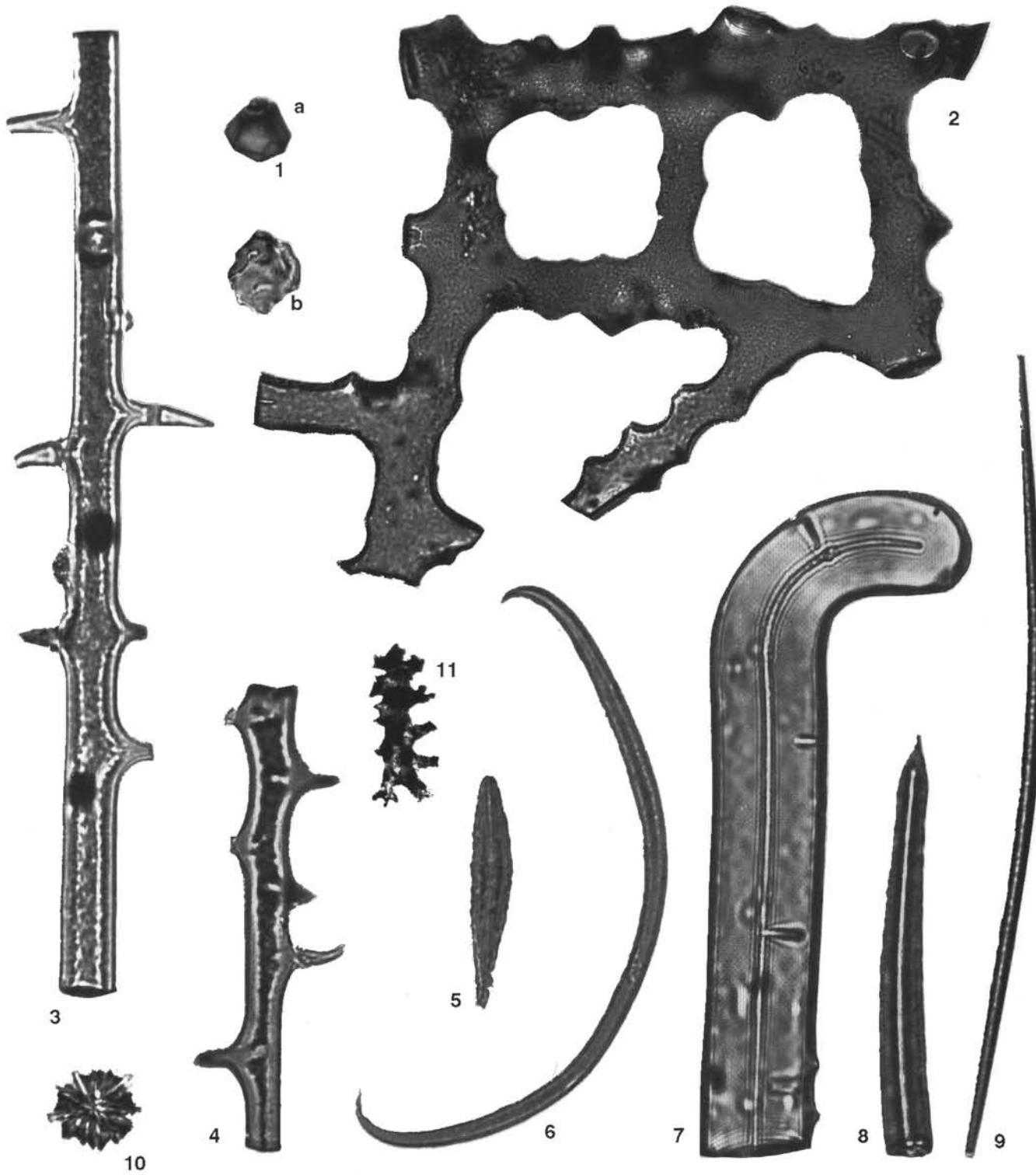


Plate 3. Magnification 1500 \times , except Figures 1b and 9 at 975 \times ; Figure 2 at 500 \times ; and Figure 11 at 300 \times . 1. Archaeomonadacea; (a) Sample 144-880A-1H-1, base; (b) Sample 144-873B-1H-2, 30–34 cm. Figures 2–9 are siliceous sponge spicules. 2. Fragment of fibroreticulate hexatinellid skeleton, Sample 144-880A-1H-1, 129–130 cm. 3–4. Spiny oxea (M). 3. Sample 144-880A-1H-1, 38 cm. 4. Sample 144-880A-1H-1, 56 cm. 5. Fusiform style (m), Sample 144-880A-1H-1, 73–74 cm. 6. Sigma (m), Sample 144-880A-1H-1, 34 cm. 7. Styloid (M), Sample 144-880A-1H-1, 2 cm. 8. Oxea? (m), Sample 144-880A-1H-1, 2 cm. 9. Microxea (m), Sample 144-880A-1H-1, 64 cm. 10. Oxyspheraster (m), Sample 144-880A-1H-2, 9–10 cm. 11. Anthosigmata (m), Sample 144-873B-1H-2, 0–2 cm.

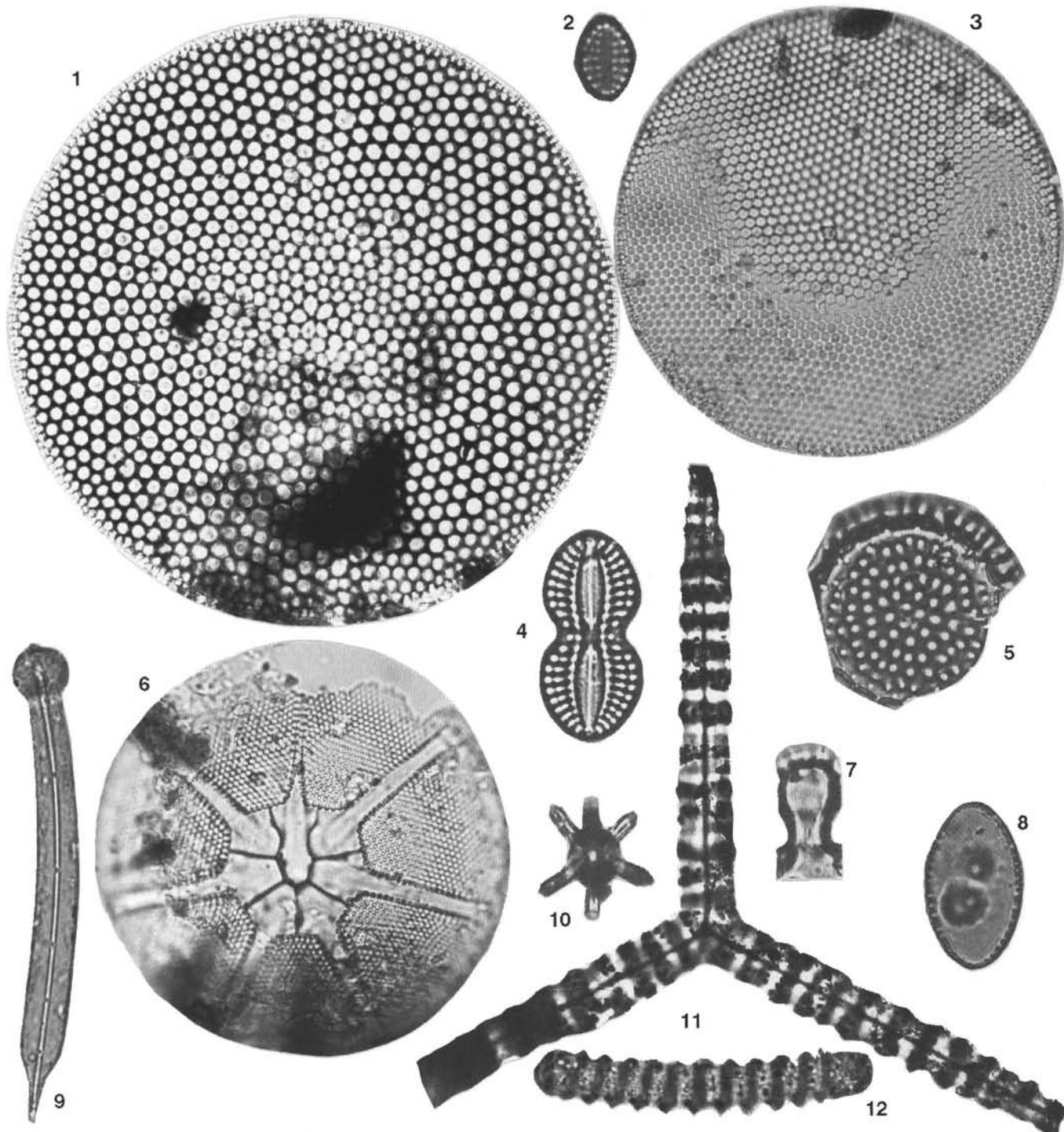


Plate 4. Magnification 1500 \times , except Figures 1, 3, 6, 9, 12 at 975 \times . Figures 1–8 are marine diatoms. 1. *Coscinodiscus pacificus*, Sample 144-878H-1R-CC. 2. *Raphoneis surirelloides*, Sample 144-880A-1H-1, 24 cm. 3. *Planktoniella sol*, Sample 144-878-1H-1 (top). 4. *Diploneis bombus*, Sample 144-880A-1H-1, 129–130 cm. 5. *Thalassiosira jouseae* (*T. nidulus*), Sample 144-880A-1H-1, 66 cm. 6. *Asteromphalus ornithopodus*, Sample 144-878A-1H-1 (top). 7. Diatom resting spore, Sample 144-880A-1H-1, 2 cm. 8. *Xanthiopyxis* sp., Sample 144-880A-1H-1, 28 cm. Figures 9–12 are siliceous sponge spicules. 9. Tylostyle (m), Sample 144-873B-1H-1, 30–34 cm. 10. Strongylaster (m), Sample 144-880A-1H-1, 64 cm. 11. Tetraxon (M), Sample 144-873B-1H-2, 0–2 cm. 12. Strongyle (m), Sample 144-873B-1H-1, 30–34 cm.